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Mapping Offshore Winds Around Iceland Using Satellite Synthetic Aperture Radar and Mesoscale Model Simulations

Charlotte Bay Hasager, Merete Badger, Nikolai Nawri, Birgitte Rugaard Furevik, Guðrún Nína Petersen, Halldór Björnsson, and Niels-Erik Clausen

Abstract—The offshore wind climate in Iceland is examined based on satellite synthetic aperture radar (SAR), coastal meteorological station measurements, and results from two atmospheric model data sets, HARMONIE and NORA10. The offshore winds in Iceland are highly influenced by the rugged coastline. Lee effects, gap flow, coastal barrier jets, and atmospheric gravity waves are not only observed in SAR, but are also modeled well from HARMONIE. Offshore meteorological observations are not available, but wind speed and wind direction measurements from coastal meteorological masts are found to compare well to nearby offshore locations observed by SAR. More than 2500 SAR scenes from the Envisat ASAR wide swath mode are used for wind energy resource estimation. The wind energy potential observed from satellite SAR shows high values above 1000 Wm\(^{-2}\) in coastal regions in the south, east, and west, with lower values in the north. The most promising region for wind energy production is the southwestern coastal region.

Index Terms—Offshore resource, polarization ratio (PR), synthetic aperture radar (SAR), wind energy, wind speed.

I. INTRODUCTION

Due to the location of Iceland both close to the North Atlantic storm tracks [1] and just east of Greenland, the wind climate of Iceland is both varied and at times extremely harsh. As wind turbines in recent decades have become more robust for applications in extreme climates, the interest of wind energy utilization in Iceland has increased. Mapping the wind resource is a vital part of understanding the feasibility. The wind conditions over Iceland have been studied for several years, using numerical models and in situ observations, in particular the distortion of the near-surface wind field by orography [2], [3], as well as a severe storm in northeast of Iceland [4]. Recently, an onshore wind atlas for Iceland was completed [5]. It confirms that the wind conditions on land are promising for wind energy production, although they are not yet exploited on a significant scale. However, there is an increased interest, and the largest power company in Iceland, Landsvirkjun, erected two wind turbines in 2013, initially for research purposes. An offshore wind resource map for Iceland would give useful additional information, if this clean energy resource is to be exploited at a later stage. Here, we present the first offshore wind resource map for the seas surrounding Iceland. The basis for most offshore wind resource maps is atmospheric modeling, the results of which are compared to offshore meteorological mast observations [6], [7]. However, offshore meteorological observations in Iceland are not available. Instead, in this study, satellite remote-sensing data are used for the estimation of wind energy resources, which has been identified as one of the nine societal benefit areas of the Global Earth Observations Systems of Systems (GEOSS) program.

The available satellite remote-sensing observations leading to surface ocean wind vectors are from microwave remote sensing. Previously, wind resource estimation has been done at a spatial resolution around 25 × 25 km in the Mediterranean Sea [8] and the Northern European Seas [9], [10], using data archives with ocean surface vector wind data from QuikSCAT. For Iceland, the most relevant areas with regard to offshore wind energy are expected to be near-coastal, i.e., less than 25 km from the coast. The coastline of Iceland is very complex in places and a spatial resolution of 25 × 25 km does not provide sufficient near-coastal coverage. Satellite synthetic aperture radar (SAR) provides microwave data useful for ocean wind mapping at around 1 × 1 km [11]–[13], which allows near-coastal wind mapping. Significant gradients in coastal winds have been observed from SAR at less complex coastlines [14], [15]. This satellite data source has therefore potential to be able to resolve winds near the coast of Iceland and is thus chosen for the study. We use SAR scenes from the European Space Agency (ESA), obtained by the Envisat satellite, which carried the advanced SAR (ASAR) instrument.

Satellite SAR was used for wind resource estimation based on a few samples from RadarSat-1 data in Canadian waters.
whereas in Japanese waters [18] and in Chinese waters [19], [20] Envisat ASAR was used. In the Northern European Seas, many Envisat ASAR samples were used for wind resource estimation [21]. Wind resource mapping is usually based on hourly values from one or more years, i.e., around 8760 observations or more. Fewer values may be used, but with increasing statistical uncertainty [22], [23].

The objective of the study is to investigate the quality of offshore winds derived from SAR in the coastal regions of Iceland.

SAR winds are compared with data from several coastal meteorological stations, as well as two atmospheric models: HARMONIE [24], [25] operated by the Icelandic Meteorological Office (IMO) and high-resolution limited area model (HIRLAM) applied in the Norwegian Reanalysis of wind and waves (NORA10) [26], [27]. The numerical models provide data on regular horizontal grids and at regular time intervals, whereas the satellite data are irregular in both space and time. However, due to the higher spatial resolution, more details in the offshore wind field can be expected from satellite observations than from modeling.

Two tests of SAR data quality will be performed.

1) Selected cases of atmospheric flow observed by SAR are examined and compared to HARMONIE model simulations, with a focus on specific meteorological phenomena in which the coastal influences on offshore winds are dominant.

2) In order to assess the overall accuracy of the applied SAR wind retrieval, a number of colocated samples of SAR and NORA10 model data are compared. Additionally, a comparison is made between coastal meteorological observations and nearby SAR offshore data.

The first test focuses on the spatial pattern of wind phenomena observed by SAR and simulated by HARMONIE, for distinct large-scale atmospheric weather patterns that are typical around Iceland. Earlier studies have compared coastal SAR winds and atmospheric model results in regions with complex coastlines near Alaska [28], [29], Svalbard [30], [31], Greenland [32], and the Philippines [33].

The second test is aimed primarily on the local agreement over time between SAR and model data. Studies from the Baltic Sea comparing SAR-derived winds to offshore meteorological wind observations from masts, dedicated specifically to wind resource assessments, show a root-mean-square error (RMSE) around 1.2 ms$^{-1}$ and a bias around $-0.25$ ms$^{-1}$ [34].

In Iceland, SAR wind maps over the open ocean are compared to NORA10 model results, as this model has been shown to successfully compared to ocean wind observations in the Norwegian Sea and the North Sea [27]. Also, a high number of colocated samples are available. More specifically, the choice of polarization ratio (PR) is determined from this analysis. The wind direction input used to drive the wind speed retrieval is from the U.S. Navy Operational Global Atmospheric Prediction System (NOGAPS) [35]. The accuracy of NOGAPS wind direction near Iceland is expected to be comparable to results from the Baltic Sea, where wind directions observed at ten offshore meteorological masts show an RMSE of 6.3°, a bias of 7.8°, and a mean squared correlation coefficient $R^2$ of 0.95 [26].

II. Study Area and Data

The study area is located between 62°–68°N and 11°–27°W. It covers Iceland and the surrounding seas. The wind observations are retrieved from seven coastal automatic weather stations; see Fig. 1 for their locations. Table I contains further information on the stations. Wind speed and direction are measured hourly at 10 m above ground level (AGL). The time period considered here ranges from 2005 to 2012. The availability of station data in percent of the dates included in the nearest SAR time series is shown in parentheses. The locations of points from which SAR time series were retrieved are indicated by blue dots.

### Table I

<table>
<thead>
<tr>
<th>Station</th>
<th>Location</th>
<th>Height (m)</th>
<th>Year</th>
</tr>
</thead>
<tbody>
<tr>
<td>Grindavik</td>
<td>63.84N, 22.417W</td>
<td>9.3</td>
<td>2008</td>
</tr>
<tr>
<td>Garðskagaviti</td>
<td>64.08N, 22.69W</td>
<td>4.0</td>
<td>1995</td>
</tr>
<tr>
<td>Gufuskálar</td>
<td>64.90N, 23.93W</td>
<td>7.0</td>
<td>1994</td>
</tr>
<tr>
<td>Grimsey</td>
<td>66.54N, 18.02W</td>
<td>19.0</td>
<td>2005</td>
</tr>
<tr>
<td>Seley</td>
<td>64.98N, 13.52W</td>
<td>18.0</td>
<td>1996</td>
</tr>
<tr>
<td>Vestmannaeyjabar</td>
<td>63.44N, 20.28W</td>
<td>40.4</td>
<td>2002</td>
</tr>
<tr>
<td>Eyrarbakki</td>
<td>63.87, 21.16W</td>
<td>3.0</td>
<td>2005</td>
</tr>
</tbody>
</table>

A. SAR Data

In total, 2581 Envisat ASAR scenes are used in the study. The number of SAR scenes per year can be found in Table II, and the number of scenes per calendar month in Table III. The number of samples per month is around 200 ($\pm 50$).
TABLE II

<table>
<thead>
<tr>
<th>Year</th>
<th>2005</th>
<th>2006</th>
<th>2007</th>
<th>2008</th>
<th>2009</th>
<th>2010</th>
<th>2011</th>
<th>2012</th>
</tr>
</thead>
<tbody>
<tr>
<td>No. of scenes</td>
<td>206</td>
<td>202</td>
<td>256</td>
<td>290</td>
<td>446</td>
<td>561</td>
<td>609</td>
<td>11</td>
</tr>
</tbody>
</table>

Note that after January 2012, the Envisat mission was terminated.

TABLE III

<table>
<thead>
<tr>
<th>Month</th>
<th>Jan</th>
<th>Feb</th>
<th>March</th>
<th>April</th>
<th>May</th>
<th>June</th>
<th>July</th>
<th>Aug</th>
<th>Sept</th>
<th>Oct</th>
<th>Nov</th>
<th>Dec</th>
</tr>
</thead>
<tbody>
<tr>
<td>No. of scenes</td>
<td>207</td>
<td>159</td>
<td>185</td>
<td>198</td>
<td>228</td>
<td>213</td>
<td>231</td>
<td>256</td>
<td>209</td>
<td>236</td>
<td>227</td>
<td>232</td>
</tr>
</tbody>
</table>

descending mode, 1339 scenes are collected within the time period of 19:53–23:52 UTC; in ascending mode, 1242 scenes between 10:58 and 13:13 UTC. SAR scenes at high latitudes are typically acquired for sea ice monitoring [36]. Horizontal polarization in transmit and receive (HH) is most suitable for this purpose, and the number of scenes in HH is thus high compared to the number of scenes with vertical polarization (VV). This data set holds a total of 2399 HH scenes and only 182 VV scenes. Fig. 2 shows the number of overlapping SAR scenes used in the study. There are more than 650 overlapping samples to the north west of Iceland, decreasing to 250 in the southeast corner of the domain. The near-shore areas of Iceland are covered by more than 400 samples.

The SAR images are calibrated and resampled to a grid cell size of 0.5 km in both horizontal directions to reduce random noise. To retrieve wind speed, it is necessary to select and apply a geophysical model function (GMF) and introduce wind direction and incidence angle to this function. The wind speed is retrieved using the GMF CMOD5.N [37] for the level at 10 m above the sea surface. It is the equivalent neutral wind speed.

The wind direction input used to drive the wind speed retrieval is from the NOGAPS [35]. The model simulations are given at 1° spatial resolution and six-hourly values. As the model has much lower spatial resolution than SAR data, the NOGAPS wind directions are interpolated spatially to each image pixel in the SAR wind retrieval [38].

Sea ice occurs in Icelandic waters during winter, mainly as frazil ice in the northern fjords. Land- and ice-covered areas are not suitable for SAR wind retrieval. A mask is therefore applied to eliminate these areas. Sea ice information is obtained from the IMS Daily Northern Hemisphere Snow and Ice Analysis at 4-km Resolution by the U.S. National Ice Center (http://nsidc.org/data/docs/noaa/g02156_ims_snow_ice_analysis/). Fig. 3 shows an example of wind speed map extending from the south of Greenland to the seas south of the
Fig. 4. Mean wind speed across the swath of the Envisat ASAR WSM scene from May 30, 2010 at 22:49 UTC (see Fig. 3). Wind speeds are given as mean values over a 50-km-wide transect and retrieved with the PR models of Thompson et al. (using $\alpha = 1.0$) and Mouche et al.

western half of Iceland. On the left panel, sea ice is clearly visible near the coast of Greenland. The right panel shows the coverage of the land and sea ice masks. The sea ice mask matches the visible ice-covered area well.

GMFs used for SAR wind retrieval are developed for data products acquired in C-band VV. A PR can be applied to adjust for the lower radar backscatter at HH polarization. Multiple formulations of the polarization exist (see the overviews by [39] and [40]). Most fall into one of two categories: 1) PRs with incidence angle dependence only and 2) PRs with incidence and azimuth angle dependence. Previous studies (e.g., [32]) indicate that the PR may also depend on the satellite sensor and calibration.

In this study, two PRs are tested on a limited set of satellite scenes prior to the wind retrieval processing for the entire data set over Iceland. The objective is to show the consequence of choosing one PR over another, more than it is to validate individual PR formulations (as offshore in situ data are not available). A subset of the Envisat ASAR scenes is processed with the formulation given originally by [41]. This model depends on the radar incidence angle and a constant called alpha. Here, the value of $\alpha$ is set to 1.0, since the original value of 0.6 is known to overestimate winds from Envisat significantly [39], [40]. The same data subset is processed with the PR model 1, developed by [39], which takes the azimuth angle into account. Collocated SAR and atmospheric model data at some distance from the coast (in open ocean) where the GMF performs at its best are selected.

B. HARMONIE Model

Additional data for the period from September 2009 to present were obtained from reanalysis project using the mesoscale model HARMONIE, version 37h1.2 [24], [25].

The model domain, with 300-by-240 horizontal grid points, and a horizontal grid-point spacing of 2.5 km in both directions, encompasses Iceland and about 100 km of the surrounding ocean. The model was run with the standard 65 vertical levels, and with a nonhydrostatic dynamic core. Radiation, turbulence,
convection, and microphysics (clouds and precipitation) were determined by the AROME upper air physics scheme. Surface and soil processes were described by the external single-layer coupled surface scheme SURFEX. Initial and lateral boundary conditions were provided by European Centre for Medium-Range Weather Forecasts (ECMWF) operational analyses. The model was run in upper air and surface data assimilation mode. The atmospheric analysis was handled such that initial and boundary conditions for each forecast run were combined with the output from the previous run (blending mode). Gridded surface analyses for 2-m air temperature and relative humidity, sea surface temperature, and snow water equivalent were prepared by the spatial interpolation tool CANARI.

C. NORA10 Model

The Norwegian Reanalysis of wind and waves (NORA10) is a downscaling of ECMWF global reanalysis ERA-40 [42] for the period 1957–2002 [26]. The downscaling is performed using the HIRLAM [43], with a 10–11-km horizontal grid-point spacing, covering the northern North Atlantic from south of Greenland to the Barents Sea and the North Sea in the east. Hourly surface winds from HIRLAM are used to run a wave model (WAM) on the same grid. From 2002 to present, the same model setup is used to downscale ECMWF forecasts. This is done to keep NORA10 updated. The model setup and downscaling procedure is described in detail by [26]. Verification of NORA10 winds is also provided therein, as well as by [27]. Note that Iceland is located close to the border of the NORA10 domain, which has an effect on the model winds in the northwest part of the area of interest.

III. RESULTS AND DISCUSSION

A. Selection of PR

The plot in Fig. 4 illustrates the difference in wind speed retrieved with the two PR models [39], [41] for a single image. The wind speed data are extracted across the entire Envisat swath with an incidence angle range of 16.6°–42.7°. Winds are averaged over the 50-km-wide transect indicated in Fig. 3. At low incidence angles, the wind speeds retrieved with the polarization of Thompson et al. [41] with \( \alpha \) set to 1.0 give much higher wind speeds than winds retrieved with the PR model of Mouche et al. [39]. The difference is more than 3 ms\(^{-1}\). The two plots approach the same value with increasing incidence angles. At 39°, the Mouche et al. model gives

Fig. 7. Envisat ASAR wind field of Iceland, observed on 15 December 2005 at 11:35 UTC. The white barbs show the NOGAPS wind directions. The gray-scale shows elevation.
higher wind speeds in agreement with previous findings of Mouche et al. [39].

Comparison of colocated ocean winds from SAR and NORA10 allows the data representation, open-ocean GMF, and PR to be tested. NORA10 and SAR winds are gridded to the same regular latitude–longitude grid over Iceland, covering $62^\circ$–$67.9^\circ$N and $24.6^\circ$–$10^\circ$W with a grid cell spacing of 0.1°. For a location in the open ocean north of Iceland (at $66.9^\circ$N, $20.7^\circ$W) with 525 collocated samples, results are shown in Fig. 5 for the two PRs. The linear correlation results have a slightly lower RMSE for Mouche et al. than for Thompson et al. We therefore choose to use Mouche et al. for this study.

Fig. 6 shows the difference between the NORA10 January 2005 and January 2012 mean wind speed and the SAR mean wind speed retrieved with the PR of Mouche et al. [39]. The upper left part in dark red (from $67^\circ$–$68^\circ$N to $22^\circ$–$24.5^\circ$W) should not be considered due to border effects at the edge of the NORA10 model domain. Near the coast, NORA10 wind speed is probably too low due to topographic effects. The difference between the two maps shows the SAR mean wind speed to be 1 ms$^{-1}$ higher near the coasts and in the northwest, while it is 0.5–1 ms$^{-1}$ lower than NORA10 in the east.

The wind speed map in Fig. 7 is from 15 December 2005 at 11:35 UTC. It is 400 km wide, covering most of Iceland.

### B. Case Studies—SAR Versus HARMONIE

Six case studies of wind conditions with significant coastal influence are examined (see Table IV). The cases are selected for different regions in Iceland. The wind field maps from SAR and HARMONIE are shown in Fig. 8. In the SAR images, the land is masked and digital elevation in gray-scale is shown. Cases B, C, and F are from the western part of Iceland, with prevailing atmospheric flow from the east and northeast. Common to all three cases are elongated streaks of elevated wind speeds through narrow fjords and over the larger bays between peninsulas, persisting for around 200 km downstream, separated by areas of lower wind speeds in the lee of coastal mountains. The wind directions from the NOGAPS and HARMONIE model compare well with the direction of the large-scale atmospheric flow features seen in SAR. Due to the higher spatial resolution, smaller scale features are visible in the SAR images more often than in the HARMONIE simulations. In general, the model

<table>
<thead>
<tr>
<th>Time</th>
<th>Key findings in wind fields SAR vs. HARMONIE</th>
</tr>
</thead>
<tbody>
<tr>
<td>A</td>
<td>Winds are from SW. Atmospheric divergent lee wake patterns east of Seley are very pronounced both in model and SAR and they extend from the coastline and more than 100 km towards the northeast. In the sea near Grímsey SAR show higher wind speed gap flow than model but the lee effects are similar in model and SAR. There is some discrepancy in wind directions.</td>
</tr>
<tr>
<td>B</td>
<td>Winds are from the east. Coastal winds are weaker in the model, especially for the gap flow in Faxaflöy, the southwestern bay, in Breiðafjörður, and the middle bay. The high winds in coastal barrier jet in the Denmark Strait near Iceland in the north are much higher in SAR than in model. The overall spatial pattern is similar between model and SAR. Model winds along the south coast compare well to those in SAR and lower winds east of the Southern Peninsula are seen both in model and SAR. Wind directions in wakes from SAR and model compare well.</td>
</tr>
<tr>
<td>C</td>
<td>Winds are from NE. The direction and extent of lee wakes west of Iceland are similar in model and SAR but SAR shows much higher wind speed compared to the model, in particular, in the Denmark Strait and higher gap flow in Faxaflöy and Breiðafjörður.</td>
</tr>
<tr>
<td>D</td>
<td>Winds are from W. The coastal lee wake off the east coast is captured well by the model in the near-shore area compared to SAR. Wind directions are similar. The wind pattern in the northeast in SAR show pronounced transverse atmospheric gravity waves while the model show less variation. The model shows clear atmospheric gravity waves from the south point of Vatnajökull and along the coastline towards the west. Also SAR shows these gravity waves whereas the wave pattern further offshore in SAR is not as regular as in the model.</td>
</tr>
<tr>
<td>E</td>
<td>Winds are from North. The pattern of wakes and northwesterly jets off the southeast coast is captured well by the model. Model winds are somewhat weaker. Wind directions are similar in SAR and model. It is clear that the SAR shows more detail in the gap flows and greater flow concentration.</td>
</tr>
<tr>
<td>F</td>
<td>Winds are from NE. Coastal lee wakes and gap flow west of Iceland from model and SAR are very similar in wind direction, wind speed and spatial pattern. The sharp gradient in wind speed south of Iceland is captured well both in model and SAR though with SAR the features are even more pronounced. In Breiðafjörður transverse lee wake pattern is seen in SAR.</td>
</tr>
</tbody>
</table>
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Fig. 8. Envisat ASAR wind speed maps with NOGAPS wind directions (left) and HARMONIE model results (right) for six cases with significant offshore coastal wind conditions in Iceland.

C. Wind Directions—SAR/NOGAPS versus Station Measurements

In order to verify SAR/NOGAPS wind directions in coastal waters, NOGAPS data are compared to measurements from onshore coastal stations around Iceland. See Fig. 1 for the locations of these stations, and the grid points from which SAR/NOGAPS wind direction time series were obtained. For each time series, the occurrence of wind direction in 45° sectors was determined (see Fig. 9). All wind direction histograms are normalized to 100%. The minimum wind speed is 3 m s$^{-1}$, which is a typical cut-in speed for large wind turbines.

As is well known, the coastal stations are significantly affected by the surrounding terrain, as well as by thermal land–sea gradients. The main problem, as far as a comparison of station and SAR/NOGAPS data is concerned, is the lack of spatial detail in NOGAPS. Since it is a global model, agreement on fine-scale wind features near the coast cannot be expected. On Seley, in the east [Fig. 9(d)], e.g., the dominant northerly winds modeled by NOGAPS are deflected toward the northeast. On Grimsey, in the north [Fig. 9(b)], the higher occurrence of southeasterly winds at the weather station, together with a reduced occurrence of northeasterly and easterly winds, is explainable by the slope of the local terrain, as the island slopes upward from west to east. At the western tip of Snæfellsnes, near Gufuskálar [Fig. 9(a)], the blocking effects from the volcano are clearly noticeable at the coast, with winds deflected either to easterly or southerly directions. These effects are absent in the wind directions from NOGAPS at the SAR reference location. However, from other directions, the wind roses are very similar. In the southwest [Fig. 9(c)], near Garðskagaviti, the blocking effects from the volcano are clearly noticeable at the coast, with winds deflected either to easterly or southerly directions. These effects are absent in the wind directions from NOGAPS at the SAR reference location. However, from other directions, the wind roses are very similar. In the southwest [Fig. 9(c)], near Garðskagaviti, the similarity between the wind direction distributions is excellent. Near Heimaey (Vestmannaeyjar), the two data sets agree about the dominant wind directions from east and southeast. However, at the nearby SAR location, the offshore wind direction distribution is more uniform than on the island. The same is true near Eyrarbakki, where the prevailing northeasterly offshore winds, affected strongly by land–sea
temperature gradients, are not found at the nearest offshore SAR location.

D. Wind Resource

Mean annual wind speed, Weibull scale (A) and shape parameters (k), and mean annual energy density at 10 m above sea level were calculated using the Satellite-WAsP (S-WAsP) [21]. The results are shown in Fig. 10. The mean wind speed ranges from 5 to 10 m s\(^{-1}\). Spatial patterns in the coastal wind field, seen for individual cases in Fig. 8, are also noticeable in the field of mean wind speed. The gap winds in the eastern fjords, the very strong winds in the Denmark Strait, and the lee wakes in Faxaflói and Breiðafjörður in the west are well-defined persistent features. Strong winds also occur along the south-western coastline, while further east, lee effects are observed near the coast.

The map of Weibull scale parameter A shows similar spatial features as the mean wind speed, though with slightly higher values. The map of Weibull shape parameter k shows relatively low values around 1.5 at the northern and eastern coastline, while k is around a more typical value of 2 along the southern and western coastlines. The Weibull shape parameter of value 1.5 indicates a wind climate with large variability in winds (flat distribution), whereas the Weibull shape parameter of value 2 indicates a wind climate with more steady winds (peaked distribution). Average energy density is between 700 and 1000 W m\(^{-2}\) along the east and southwest coast. However, the highest values in these regions are found very close to the coast, and may be artifacts of image processing. The lowest values of 500–700 W m\(^{-2}\) occur along the north and southeast coast. Along most of the western coastline, the energy density is low as well, with the exception of most of the Westfjords. There, the highest values of 1400 W m\(^{-2}\) very close to the land are most likely also artifacts of image processing. The wind power is proportional to the cube of the wind speed, so error in SAR or model data will result in high error on wind power.

The high SAR wind speeds in the Denmark Strait may be affected by ocean currents. However, this region is not likely to be a choice for wind energy utilization, and will not be considered here.

For comparison, mean annual wind speed and power density, as well as the annual values of Weibull scale and shape parameters, were calculated based on HARMONIE reanalysis data (see Fig. 11). Neither the coverage during the day nor the time periods are the same for the SAR measurements and the HARMONIE reanalyses. However, the values in the coastal zone around Iceland are very similar for the two data sets. As for the SAR data, the lowest HARMONIE wind speeds of around 5–7 m s\(^{-1}\) occur along the north and southeast coast, and the highest wind speeds of around 10 m s\(^{-1}\) are found at the exposed tips of peninsulas in the west. The Weibull shape parameter also has a similar pattern as the average wind speed, but equally with slightly higher values. As for the SAR data, low values of power density of 500 W m\(^{-2}\) occur along the north and southeast coast. However, the high SAR values of
1000 Wm$^{-2}$ close to the east and southwest coast are not found in the model, suggesting again that they might be artifacts or SAR can resolve features smaller than the model. Based on HARMONIE data, the Weibull shape parameter is close to 2 along most of the Icelandic coastline. As for the SAR data, the lowest values exist off the northeast coast of the Westfjords. However, overall, the SAR-derived values are lower along the north and east coast, indicating that temporal variability of wind speed based on the satellite measurements is higher than in the mesoscale model.

E. Discussion

When satellite wind fields are used for offshore wind resource estimation in a new geographical area, it is advisable to critically investigate a set of issues. The first issue is the satellite wind field processing accuracy. The second issue is the spatial variability of observed winds. The third issue is the statistical representation of satellite wind field maps, including temporal variability.

Satellite SAR is expected to be useful for offshore wind resource estimation, with the capability to quantify coastal wind phenomena around Iceland at finer scales than through numerical modeling. The lack of offshore meteorological observations resulted in the following choices for the investigation of the satellite SAR wind fields.

The choice of GMF is CMOD5.N, one of the most recent models, and valid up to hurricane strength winds. The Envisat ASAR archive for Iceland is dominated by HH polarization scenes, and the selection of PR is thus important. We use NORA10, collocated to SAR, to select the PR. The statistics between the tested PRs were not very different though. Further investigation to assess the consequences of the choice of PR for the mean wind speed map could be done. Testing of NOGAPS wind directions was performed by comparing wind roses observed at coastal weather stations in Iceland and nearby SAR/NOGAPS wind roses. The local effects of rugged terrain influence the coastal wind roses considerably, while the resolution of NOGAPS does not provide fine-scale detail. However, the agreement among most sites is good. The discrepancy in mean speed for collocated SAR and NORA10 samples is unexplained for the region east of Iceland, where NORA10 shows 0.5–1.0 ms$^{-1}$ higher values than SAR. In most other areas, the discrepancy is lower. The underestimation of winds from NORA10 near coastlines and in the northwest area is explained by the model resolution and domain setup, respectively. It is clear that further investigation of PR, GMF, and wind direction near Iceland would be useful, in particular, if in situ observations become available, but even then it would be difficult to separate these effects. The recent work of [40] could be an option for improvement on PR.

The spatial variability of SAR wind speeds is investigated for six selected cases, using HARMONIE for comparison. Enhanced gap flow and lee-wake effects occur for all offshore wind directions around Iceland. For most cases, the wind directions between the HARMONIE and NOGAPS models agree.
well. All reported cases have medium-to-high winds (around 6–20 m s\(^{-1}\)). For very low wind speeds (less than 3 m s\(^{-1}\)), wind directions are more variable and have relative higher uncertainty. The SAR-derived wind maps at 1 × 1 km resolution have much spatial detail. According to [44], the spectral density in Envisat ASAR WSM wind maps is high. Thus, SAR wind fields observe small-scale variability of ocean surface winds. The smallest of these scales would not be modeled by HARMONIE, but the overall agreement between SAR and HARMONIE is good. Sharp gradients in wind speed are seen in Cases B, C, and F. It is interesting to notice the presence of atmospheric gravity waves both in SAR and HARMONIE. These, among other coastal wind phenomena, such as barrier jets, are also discussed in [28], [29], and [45]. The good agreement between SAR and HARMONIE model, regarding the representation of significant spatial patterns in coastal winds around Iceland, supports the combined use of SAR-derived wind maps and mesoscale model simulations for wind resource estimations.

The third issue to investigate is the statistical representation of satellite wind field maps, including temporal variability of the data set for wind resource statistics. The number of overlapping SAR wind maps near the coast of Iceland is more than 400, which is expected to be sufficient for prefeasibility wind resource estimation [46]. The representation of seasonal variability in winds seems to be reasonably well captured in the present data set with monthly numbers of samples of around 200 (±50). The diurnal sampling is dictated by the satellite orbit, and SAR winds are observed only during two 3-h periods around noon and late evening. However, as shown in the Section III-E in comparison with the HARMONIE results, the SAR-derived wind field represents well the annual wind conditions.

The most promising coastal regions for wind energy production are along the southwestern coastline, with a mean annual energy density of around 700–1000 W m\(^{-2}\), and with values of the Weibull shape parameter \(k\) around 2. Along the northern coast, the energy density in several areas is above 1200 W m\(^{-2}\). However, this may be an artifact, due to sea ice not being fully avoided, despite the sea ice mask being applied. Areas with sea ice tend to give overestimated wind speed. This is also true around small peninsulas and islands, which are assumed to be water, but in reality are hard targets.

IV. CONCLUSION

The wind resource results from satellite SAR, combined with HARMONIE model simulations, are useful for prefeasibility studies of the offshore wind energy resource around Iceland, i.e., early planning at a strategic level, while further investigation is needed for siting and development of wind farms.

The satellite SAR data indicate a high wind energy resource above 1000 W m\(^{-2}\) along the southwest coast, with slightly lower values based on model simulations. These results need further validation from on-site measurements, in case siting and
development of wind farms is to take place. The limitations in the current assessment of winds from SAR around Iceland are that the selected GMF, PR, and wind direction input could not be verified with direct offshore observations.

Several specific atmospheric phenomena are captured well by both SAR and HARMONIE, with greater detail based on the satellite measurements. However, despite the lower resolution, the mesoscale model also has certain advantages. (a) Results are not affected by the presence of sea ice. (b) Data are regular (unbiased) in space and time. (c) Data are available at higher levels above the sea surface. (d) Wind direction is spatially color-coded with wind speed, with a higher resolution than that from NOGAPS. A combination of the two data sources therefore gives the best results.

REFERENCES

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