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Timing of the compensation of winter respiratory carbon losses provides explanatory power for net ecosystem productivity of forests

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Key words: eddy covariance; CO₂ exchange; carbon sink; carbon source; growing season length; net ecosystem productivity; NEP prediction; winter respiration; forest ecosystems

Key Points (max 100 characters per item):

- Compensation day (cDOY) is the day of year when net C losses during winter are compensated by net C uptake in spring
- cDOY largely explains annual net ecosystem productivity NEP_c of forests when the site has a distinct winter respiratory loss period
- cDOY and its explanatory power depends on the integration method for annual NEP_c and on the forest type
- The findings highlight the importance of carry-over effects and winter time processes to understand NEP_c

Abstract

Accurate predictions of net ecosystem productivity (NEP_c) of forest ecosystems are essential for climate change decisions and requirements in the context of national forest growth and greenhouse gas inventories. However, drivers and underlying mechanisms determining NEP_c (e.g. climate, nutrients) are not entirely understood yet, particularly when considering the influence of past periods.

Here we explored the explanatory power of the compensation day (cDOY) —defined as the day of year when winter net carbon losses are compensated by spring assimilation— for NEP_c in 26 forests in Europe, North America, and Australia, using different NEP_c integration methods.

We found cDOY to be a particularly powerful predictor for NEP_c of temperate evergreen needle-leaf forests ($R^2 = 0.58$) and deciduous broadleaf forests ($R^2 = 0.68$). In general, the latest cDOY correlated with the lowest NEP_c . The explanatory power of cDOY depended on the integration method for NEP_c , forest type, and whether the site had a distinct winter net respiratory carbon loss or not. The integration methods starting in autumn led to better predictions of NEP_c from cDOY than the classical calendar method starting at January 1. Limited explanatory power of cDOY for NEP_c was found for warmer sites with no distinct winter respiratory loss period.

Our findings highlight the importance of the influence of winter processes and the delayed responses of previous seasons' climatic conditions on current year's NEP_c . Such carry-over effects may contain information from climatic conditions, carbon storage levels and hydraulic traits of several years back in time.

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1 Introduction

Accurate predictions of carbon dioxide (CO₂) exchange by forest ecosystems are essential for understanding, e.g., the role of the forest mitigation in the context of the National Determined Contribution under the Paris Agreement, as well as for the required estimates of annual carbon (C) budgets to be provided at national or international level. Research in the past decades focused on improving these predictions on both annual and longer (decadal) time scales, e.g. in relation to extreme events (e.g. *D Baldocchi and Wilson* [2001]; *Ciais et al.* [2005]; *Richardson et al.* [2009]; *Rodrigues et al.* [2011]; *Wolf et al.* [2013]; *C Wu et al.* [2013]), and in relation to the length of the growing seasons or the number of carbon uptake days (e.g. *Churkina et al.* [2005]). Our study builds on the current understanding that some critical periods within the current or the past year (e.g. winter frost, spring drought) may explain the inter-annual variability of C uptake of forests better than average conditions over the current year only [*Le Maire et al.*, 2010]. The effect of climatic conditions from previous seasonal periods on current year's annual net ecosystem productivity (NEP_c) are called carry-over effects and were quantified e.g. by *Shao et al.* [2016]; *Thomas et al.* [2009]; *Zielis et al.* [2014]. Such carry-over effects support the influence of specific periods in the past on current year's NEP_c, and their influence have been demonstrated a long time ago by tree-ring analyses, e.g., for Danish forests [*Holmsgaard*, 1955]. Here, we use positive NEP is defined as net C uptake, while negative NEP is a net C release to the atmosphere (NEE is defined with the opposite sign in *Aubinet et al.* [2012]). Further, NEP_c is defined as the cumulative sum of NEP fluxes throughout the annual cycle – not necessarily a calendar year – yielding net C flux between the atmosphere and the forest. To account for the temporal integration of the average NEP_c over an annual cycle, values are expressed in g C m⁻² yr⁻¹, whereas half-hourly NEP measurements are given in g C m⁻² s⁻¹ (as calculated from μmol m⁻² s⁻¹).

1.1 The concept of cDOY

Following the concept of previous year's weather conditions influencing current year NEP_c, we explored the information content of cDOY, defined as the day of year when the net carbon losses accumulated during the wintertime are compensated by net assimilation in spring. The timing of cDOY is assumed to change with climatic conditions of previous periods (of unknown length) and may have a direct impact on the current year NEP_c [*Zielis et*

al., 2014]. Similar approaches were described in literature, e.g. the ‘zero-crossing time’, wherein net ecosystem exchange is used to define the time when the forest ecosystem turns from a C-source in winter into a C-sink in spring [Gonsamo *et al.*, 2012a; Gonsamo *et al.*, 2012b]. Another approach quantifies the so-called ‘start of the carbon uptake period’ (CUP) which is determined by a sharp increase in gross primary production (GPP) [Delpierre *et al.*, 2009]. However, these approaches rely on instant net ecosystem exchange rates only, and do not accumulate carbon loss over an entire wintertime, as it is the case of in the cDOY approach.

1.2 Integration methods for NEP_c

Traditionally, NEP_c is integrated annually over a time period of the Gregorian calendar year (classical integration as shown in Fig. 1). This is more a practical choice, but it neither reflects any particular connection to underlying carbon cycle processes, nor does it take into account potential carry-over effects on NEP_c . As an example, trees prepare their buds in autumn and thus the predisposition for growth (and thus NEP_c) during the following season is determined in autumn already. Thus, it is important to consider the start and end of the accumulation period of NEP_c . In line with these thoughts, Urbanski *et al.* [2007] introduced a method integrating NEP_c at Harvard Forest from October 28 to October 27 of the following year (Urbanski integration in Fig. 1), trying to come closer to a more reasonable biological time reference of the annual NEP cycle. This integration period is similar to the hydrological year as starting on November 1 in the northern hemisphere. Thomas *et al.* [2009] found that inter-annual and seasonal variations in carbon and water processes were best explained when seasonality was defined functionally within hydrological years.

More recently, a dynamic integration approach was introduced by Zweifel *et al.* [2010] in order to relate continuous stem diameter fluctuations to NEP_c . The frequently occurring stem shrinkages induced by winter frost [Zweifel and Hasler, 2000] made the classical integration approach from 1st of January to 31st of December inapplicable for an unbiased analysis of annual stem growth increments (bark and wood) in relation to NEP_c . An integration over a variable period was therefore proposed (Fig. 1), starting with the day when NEP_c of the previous calendar year reached its maximum and ending with the day in the current year when maximum NEP_c was achieved (dynamic integration in Fig. 1). Thus, the dynamic year

corresponds more closely to the actual biological cycle, which does not exactly count 365.25 days per year. This dynamic integration method is appropriate to time series of stem increments and NEP_c data from eddy covariance flux measurements. It was concluded that the application of this approach reduces distortion effects on annual sums, due to apparent inter-annual variations in carbon losses and stem shrinkages during wintertime [Zweifel *et al.*, 2010], i.e., shifts in uptake and loss periods that arbitrarily affect the sums calculated over a fixed calendar period. We use the terms ‘year’ and ‘annual’ in combination with all three integration methods, for the sake of readability, being aware that the terms usually are implicitly used for periods of Gregorian calendar years from 1st of January to 31st of December.

The way of splitting time series into annual integration periods also changes the potential contribution of winter carbon losses for the total annual C uptake and the cDOY timing in the following year (Fig. 1). Indeed, the classical integration period splits the net carbon loss of a winter period in two parts, assigning them to two different NEP_c years, whilst the dynamic (and Urbanski) integration method assigns net carbon loss for all the winter period entirely to the NEP_c of the biological year that will last until the onset of the next winter period. Accordingly, cDOY changes with the respective integration method (Fig. 1) and might have a different explanatory power for NEP_c .

In this study, we used in total 26 eddy covariance forest sites with 25 sites throughout Europe, and North America (Fig. 2), and additionally one site from Australia, thus covering a wide range of climatic conditions (Table 1) to investigate the meaning of cDOY for NEP_c and its underlying drivers. We used the cDOY timing as the key measure associated with the net carbon loss period and related it to climatic conditions and NEP_c . We addressed the following specific objectives: (1) application of three different NEP_c integration methods (classical, Urbanski, and dynamic) in order to calculate and compare the respective cDOYs; (2) identification of climatic and biological drivers for cDOY across sites and across different years; (3) evaluation of different cDOY as a predictor for its associated NEP_c ; and (4) the weight of winter net respiratory losses on current year’s NEP_c .

2 Materials and Methods

2.1 Study sites

The study is based on carbon dioxide (CO₂) flux data from 347 site-years from 26 eddy covariance (EC) forest sites ('managed forest not affected by major disturbances like fire or wind throw) within Europe, North America, and Australia (Table 1, Fig. 2). The selected sites fulfilled the following criteria: (1) at least four years of continuous EC data; (2) availability of Level 4 (L4) data quality according to the European Fluxes Database [*European Fluxes Database Cluster*, 2014] or available from the FluxNet2015 dataset (<http://fluxnet.fluxdata.org/data/fluxnet2015-dataset/>); and (3) available meteorological and forest characteristics data. The forest vegetation at the sites was classified as deciduous broad-leaf forests (DBF, n=7), mixed forest (MF, n=3), evergreen needle-leaf forests (ENF, n=13), and evergreen broad-leaved forest (EBF, n=3).

2.2 CO₂ flux measurements

Half-hourly or hourly CO₂ flux data (net ecosystem exchange rates summed up to net ecosystem productivity, NEP), derived from both open- and closed-path gas analyzers were downloaded from the FluxNet2015 dataset (<http://fluxnet.fluxdata.org/data/fluxnet2015-dataset/>) or in L4 quality from the European Fluxes Database (<http://www.europe-fluxdata.eu/home>). These data were already filtered and gap-filled (Table 1). For three sites (CH-DAV, CH-LAE, PL-TUC, all open-path gas analyzers) our own site-specific processing was conducted: data were filtered for unfavorable atmospheric conditions such as snow, heavy rain and/or dust which increased window dirtiness of the infra-red gas analyzer > 70%. For these three individual sites, the threshold for insufficient nocturnal turbulent mixing of the atmosphere (determined via the friction velocity u_* for mechanical turbulence) was determined with the online EC gap-filling and flux partitioning tool (Markus Reichstein, Olaf Menzer, <http://www.bgc-jena.mpg.de/~MDIwork/eddyproc/>) [Reichstein *et al.*, 2013; Reichstein *et al.*, 2005] and was found to be 0.2 m s⁻¹.

2.3 CO₂ flux integration

Annual NEP_c was integrated with three different methods (Fig. 1): classical method: NEP_c is integrated from January 1 to December 31; Urbanski method: NEP_c is integrated from October 28 to October 27 one year later [Urbanski *et al.*, 2007]; dynamic method: integration of NEP_c from the DOY with the maximum seasonal peak of the previous year (typically in fall; MAX_{NEP_c}) to the DOY with MAX_{NEP_c} of the current year (Fig. 1). The dynamic integration method led to ‘annual cycles’ ranging from 7 to 16 months depending on year and site; the overall average was 364 days (Supplement Fig. S2). For the Southern hemisphere site AU-Tum, the year has been shifted half a year forward, i.e. the classical year started with July 1 and the dynamic ‘integration period’ started with the maximum peak (MAX_{NEP_c}) before July 1. The Urbanski integration method was not applied for this site.

2.4 Statistical analyses

Statistical analyses were performed using the statistical software R, version 3.3.1 [R Development Core Team, 2013]. All multiple regression models were based on linear relationships. Adjusted R² (adjR²) was used for the quantification of goodness of fit. Analyzed potential drivers for cDOY are listed in Table 1. Their respective impacts on cDOY were analyzed with multiple regression models based on the inclusion of explaining variables in a stepwise way. The so-called standardized regression coefficients (β -coefficients) were used to determine the relative importance of variables (var) within the models, ranging between -1 as the highest negative and +1 as the highest positive correlative importance [Quinn and Keough, 2002]. A β -coefficient close to zero indicates that the variable does not add to the quality of the model.

3 Results

3.1 Compensation of net carbon loss after wintertime

The day of year when respiratory carbon losses from the previous winter were compensated (cDOY) differed strongly across sites (Table 2, Fig. 3). cDOY varied from January 3 (AU-TUM/July 3) to July 25 (CA-QFO), with a mean of May 3 (obtained by averaging all three integration methods, Fig. 3). Some sites showed no or irregular cDOY timings, meaning that they observed no distinct respiratory carbon loss period every year (Table 2). Evergreen forests (3xEBF, 13xENF) in general had an earlier cDOY (April 18) than deciduous forests (7xDBF, June 28). Only nine out of the 26 sites compensated on average their net carbon losses in the climatologically defined spring calendar months (Mar–May) (Table 2). Six sites compensated before spring, while eleven compensated after May. The yearly standard deviation of cDOY for individual sites ranged from six days (DE-HAI) to more than 50 days (PT-ESP) (Table 2).

Further, cDOY strongly depended on the integration method. In general, the classical integration method led to a cDOY almost three weeks earlier than those obtained with the dynamic method (classical: April 16, dynamic: May 10). The average cDOY obtained from the Urbanski integration method (May 5) was almost the same as that from the dynamic method (data now shown). Much less affected were the mean differences (Urbanski vs. classical: $-54 \text{ g C m}^{-2} \text{ yr}^{-1}$, dynamic vs. classical: $-91 \text{ g C m}^{-2} \text{ yr}^{-1}$) and the standard deviations (Urbanski vs. classical: $-7 \text{ g C m}^{-2} \text{ yr}^{-1}$, dynamic vs. classical: $-10 \text{ g C m}^{-2} \text{ yr}^{-1}$) of NEP_c between the different integration methods (see also Supplementary Figs for each site).

3.2 Drivers of cDOY

Average cDOY was substantially correlated with mean annual air temperature (R^2 between 0.4 and 0.45). The relationship was largely independent of the integration method used (Table 3) and the later cDOYs corresponded to the cooler sites (Fig. 3a). Other site characteristics considered (latitude, longitude, altitude, tree age, nitrogen deposition, tree height, and mean annual precipitation) showed weak (or no) linear relationship to cDOY and did not improve the stepwise multiple linear regression models to explain cDOY (Table 3).

The meaning of mean annual site temperature (MAT) for cDOY was markedly increased when the pooled data over all sites were grouped into the four forest types (Fig. 3a): evergreen needle-leaf forest (ENF, all included), evergreen broad-leaf forest (EBF), mixed forest (MF) showed R^2 between 0.64 and 0.99. No significant correlation was found between MAT and cDOY for the deciduous broad-leaf forests (DBF; $R^2 = 0.07$, $p > 0.05$).

In Fig. 3b, those sites without a distinct winter respiratory loss period, and thus with no consistent cDOY timing (Table 2) were removed (all EBF and more than 50% of the ENF sites). All of these sites are evergreen, with a majority having MAT over 8-10°C, hence, in winter, these sites likely photosynthesize. The remaining six ENF sites (CA-QFO, CH-DAV, FI-HYY, RU-FYO, SE-NOR, US-NR1) with a distinct winter respiratory loss and a latter cDOY, increases to an $\text{adj}R^2$ of 0.90 for the linear relationship between MAT and cDOY (Fig. 3b).

3.3 Relationship between cDOY and NEP_c

The 26 sites analyzed in this study included C sink and C source sites (Table 2). The largest net annual respiratory loss was at RU-FYO with a consistent average C output of 137 g C m⁻² yr⁻¹. The largest net C uptake was at AU-Tum with 1007 g C m⁻² yr⁻¹.

Stepwise multivariate analysis showed that cDOY, among the site characteristic variables available, explained most of NEP_c for all integration methods (Table 4). Sites with distinct winter respiratory loss, explained significantly more of NEP_c than all other sites. cDOY obtained from the two integration approaches that initiated the NEP_c -year in autumn (Urbanski and dynamic) explained NEP_c significantly better ($\text{adj}R^2 = 0.35$ and 0.47) than cDOY from the classical integration approach ($\text{adj}R^2 = 0.23$). When the ENF (Table 4d) and DBF (Table 4e) sites were analyzed separately (using the dynamic integration), the R^2 of the linear regressions was further improved (R^2 of 0.58 and 0.68, respectively).

Mean annual temperature (MAT) was the secondary determinant variable of NEP_c in stepwise multiple linear regression models (Table 4). The ranking of site factors, with minor contributions, such as leaf area index (LAI), mean annual precipitation (MAP) and stand age followed next, however the ranking depended on the integration method. An exception was the DBF sites (Table 4e): MAT had no explanatory weight for NEP_c at these sites, in line with the finding that cDOY of these forests were not determined by MAT (Fig. 3).

When analyzing individual sites instead of pooled data, the site-specific relationships between cDOY and NEP_c showed a high variability and ranged from not existing to excellent (annual resolution, Tab. 1, Fig. 4, Fig. 5). There appeared clear clusters of points (in the scatterplot of cDOY vs. NEP_c , according to the forest types in Fig. 4): evergreen forests (ENF and EBF) had the lowest cDOY with the highest NEP_c . Deciduous broad-leaf forest (DBF) had the highest cDOYs with on average lower NEP_c . Mixed forest (MF) had average cDOYs with relatively high NEP_c .

The site-specific quality of the relationships between cDOY and NEP_c was largely explicable by grouping the pooled data according to sites with and without a distinct net respiratory carbon loss over wintertime (Table 2). The separation criterion for the two groups was a net respiration loss of 10% of the annual NEP_c (Table 2). Sites with distinct winter respiration loss had on average a stronger correlation between cDOY and NEP_c (R^2 0.53 vs. 0.37; dynamic method), and were on average 4°C cooler than sites with no distinct winter respiration loss (Tables 1 and 2).

3.4 Quality of NEP_c predictions from cDOY

The quality of NEP_c predictions from cDOY were tested by comparing measured and modelled NEP_c per site and year with a leave-one-year-out cross-evaluation (Fig. 5). There were two very clear results: (i) the Urbanski and the dynamic integration methods led to distinctly better NEP_c prediction than the classical integration method over the Gregorian/orbital calendar year. And; (ii) NEP_c predictions from cDOY were stronger for sites with a distinct respiratory carbon loss over wintertime. Thereby, sites where the forest did not become a C-source for a distinct period, and thus did not lose at least 10% of annual NEP_c every year, failed to show a strong prediction of NEP_c from cDOY.

4 Discussion

There is increasing evidence that a considerable proportion of the inter-annual variability of NEP_c cannot be explained by the current year's climatic variability alone, but needs considering previous periods' weather factors. Predisposition of growth by the determination of buds in autumn of the past year [Thomas *et al.*, 2009; Zweifel *et al.*, 2006], carry-over effects on physiology in years following climate extremes [Law *et al.*, 2002; Thomas *et al.*, 2009; J Wu *et al.*, 2012; Zielis *et al.*, 2014], C-storage pools accumulated over several years [Carnioli *et al.*, 2009; Hoch *et al.*, 2003], sapwood related hydraulic traits [Zweifel *et al.*, 2006], and winter chilling effects (vernalization) [B I Cook *et al.*, 2012] are examples of potential causalities between conditions back in time and the current-year NEP_c . In order to better understand the intra-annual variability of NEP_c and its drivers, we introduced the day of compensation (cDOY), i.e., the day of the current year (typically in spring) when net carbon losses during wintertime are compensated by carbon assimilation in spring or early summer (Fig. 1). cDOY reflects the complete winter conditions and the related accumulated CO_2 losses, in combination with the onset and rate of CO_2 assimilation in spring (Table 2). Therefore, cDOY is not directly comparable with studies focusing on the onset of GPP or the change in NEP/NEE from a C-source to a sink in spring [Delpierre *et al.*, 2009; Gonsamo *et al.*, 2012a; Gonsamo *et al.*, 2012b] since these approaches do not account for the amount of accumulated respiratory C losses over wintertime. In the following we discuss the meaning of cDOY and its impact on the interpretation of NEP_c .

4.1 Mean annual site temperature determining cDOY

The loss of C during wintertime and the respective cDOY was found to be statistically highly independent of most of the site characteristics like mean annual precipitation, nitrogen deposition, leaf area index, age, or tree height (Table 3). Only MAT was significantly related to cDOY (R^2 about 0.4, pooled data for all sites, Table 3) particularly when the sites were grouped according to their forest types (R^2 up to 0.99, Fig. 3, with one exception, see below).

The importance of air and soil temperatures for the recovery of trees from the inactive physiological winter dormancy back into a physiologically active status is well documented [D D Baldocchi *et al.*, 2005] and covers issues such as rehydration of tissues [Koike, 1990; Lundmark *et al.*, 1988; Suni *et al.*, 2003; Zweifel *et al.*, 2000]), bud burst [Basler and Körner,

2014], assimilation [Monson *et al.*, 2011a], flowering [B I Cook *et al.*, 2012], length of the vegetation/growth period [Aurela *et al.*, 2004; D Baldocchi and Wilson, 2001; Churkina *et al.*, 2005; Monson *et al.*, 2011a], growth [Zweifel *et al.*, 2010], and probably many more. All these processes are, finally, determining cDOY with different weights, since they are influencing quantities and timing of ecosystem respiration and assimilation, explaining the influence of MAT on cDOY well.

4.1.1 One exception: the deciduous broad-leaf forests

There was one exception from the generally close relationship between MAT and cDOY: MAT had no impact on cDOY for deciduous broad-leaf forests (DBF, $n=7$) (Fig. 3) but cDOY had a high explanatory power for NEP_c, particularly with the dynamic integration method (Table 4e). This finding was unchanged when considering DBF filtering for those sites with a distinct winter respiratory carbon loss of more than 10% of the annual NEP_c (Fig. 3b, negative sign = respiratory loss/positive NEP_c). Overall, this means that cDOY is strongly forest-type specific and that cDOY includes information not covered by the site characteristics investigated and thus offers a new dimension in interpreting NEP_c. This seems to be particularly true for the DBF sites. The seven DBF forests included in this study (IT-COL, US-MMS, DE-HAI, US-HA1, US-WCR, DK-SOR, US-UMB) consisted of beech (*Fagus sylvatica*), maple (*Acer spp.*), oak (*Quercus spp.*), ash (*Fraxinus spp.*), basswood (*Tilia americana*) and sourwood trees (*Oxydendrum arboreum*). We suggest two potential explanations why the cDOY of these forests does not depend on MAT. First, (i) the group of DBF sites might still be too heterogeneous in terms of their species composition to show a concise MAT-cDOY relationship. The limited number of replications ($n=7$) for this group does, however, not allow for further differentiations. And second, (ii) cDOY reflects processes which are indeed independent of MAT for this forest type, e.g., due to biological predispositions of water and carbon storage which have their origin before the time period investigated [Keenan *et al.*, 2012; Urbanski *et al.*, 2007; Zielis *et al.*, 2014; Zweifel *et al.*, 2010], or due to genetic predispositions which determine the regulation of physiological activity independently of temperature [Basler and Körner, 2014], or in a way that positive and negative temperature effects level each other off. A convincing chain of arguments for the second explanation was recently brought up by B I Cook *et al.* [2012]. They showed that increasing temperatures during winter and spring induce opposite effects in certain species.

Warmer winter conditions can lead to an insufficient vernalization, i.e., chilling requirements that must be met before a plant is able to respond to spring warming, which in turn leads to a delayed initiation of phenological processes in spring despite the positive effect of increased spring temperatures. Further, for beech trees [Basler and Körner, 2014] recently reported a co-determination of beech bud burst by the photoperiod, and, therefore, a partial decoupling from temperature. Such a partial decoupling from temperature in terms of physiological processes could be, in terms of physiological processes, a species-specific explanation for a predisposition disturbing the generally valid relationship between MAT and cDOY. The effect of climate change on the relationship between cDOY and NEP might thus also depend on species-specific physiological responses and acclimation potentials. It is however difficult to understand how heterotrophic respiration (R_H) in the soil is triggered by the mentioned tree physiological processes. Apart from temperature, R_H might be stimulated by rhizosphere processes such as root exudates and mycorrhiza, which in turn might be more closely coupled to the tree physiological status in DBF. Further field studies are needed to test this hypothesis.

4.3 Timing of cDOY

In general, evergreen forests (EBF and ENF) had earlier cDOYs than the deciduous forests, and mixed forests with evergreen and deciduous species were in between (Fig. 4). Photosynthesis of evergreens during winter varies with climatic region, but can be substantial. Thus, the early cDOYs of the evergreens may be explained by the ability of evergreen trees to start earlier in the season with assimilation [Richardson *et al.*, 2010] or even maintain it during mild winters [Pallardy, 2010]. Photosynthetic capacity can be attained after just a few days of sufficient environmental conditions [Ottander *et al.*, 1995; Ottander and Oquist, 1991; Suni *et al.*, 2003].

Forest types, excepting DBF, and cDOY are both found to be linked to MAT (Fig. 3, Table 3). Evergreen broad-leaf forest (EBF), for instance, grow at relatively warm sites, and do not have a consistently occurring winter respiratory carbon loss period and thus show no consistently cDOY timings in each year (Fig. 3). Typical examples are the eucalypt sites in Australia and Portugal (Table 2, Supplementary Fig. S AU-TUM and S PT-ESP). These sites show an almost full year growth period or at least do not turn into C-sources once every year,

and the cDOYs, which can hardly occur, happen thus hardly any year. At the other end of the biological scale appear the deciduous broad-leaf tree sites (DBF) with the latest cDOYs (Fig. 4) at the generally cooler sites (Fig. 3). The existence of a cold season is the main reason for forming a deciduous canopy. Deciduous forests need more time in spring for bud burst and leaf flushing, for the development of the photosynthetic apparatus, and the onset of photosynthetic activity [Basler and Körner, 2014; Epron *et al.*, 1996; Jurik, 1986; Koike, 1990; Reich *et al.*, 1991]. The evergreen needle-leaf forests (ENF) have the widest temporal range for cDOY (Fig. 4), again in line with the widest range of occurring MAT (Fig. 3a).

4.4 Strengths and limitations of cDOY to predict NEP_c

The explanatory power of cDOY as a predictor of NEP_c was strongly depending on whether the site had a net carbon respiratory loss higher than 10% of the annual NEP_c or not (Figs. 4 and 5). For sites with a distinct net carbon loss over wintertime (Fig. 5a) the estimated annual NEP_c from cDOY reached accuracies of $\pm 75 \text{ g C m}^{-2} \text{ yr}^{-1}$ which are comparable to some of the most successful (but much more complex) NEP models [Keenan *et al.*, 2012]. For the other sites without a distinct winter respiratory loss, the standard deviation between modeled and measured NEP_c was a factor 2 to 3 higher (Fig. 5b), which leads to the conclusion that cDOY is of limited explanatory power in these cases. This could be explained partly by the large variation in winter photosynthesis in temperate evergreens, and by the fact that evergreen needle-leaf species grow in some of the harshest conditions, such as the western US where summer drought is the norm [Law and Waring, 2015].

Besides the importance of the winter net respiratory C loss, the forest type had a strong influence on the predictive power of cDOY on NEP_c . Pooled data reached an R^2 of 0.47 (dynamic integration) for the linear regression between cDOY and NEP_c (Table 4), whereas the grouped data for ENF ($R^2 = 0.58$, dynamic integration, Table 4e) and DBF ($R^2 = 0.68$, dynamic integration, Table 4d) were much higher. This again indicates that the information content of cDOY, i.e., the net effect of winter and spring processes, depends on the forest type and the respective species composition (Fig. 4). Both winter respiratory loss and vegetation type are related to temperature and therefore linked to each other (Fig. 3). It is therefore not surprising that besides cDOY as the variable with the highest explanatory power for NEP_c , mean annual temperature appeared as the second best driver in our stepwise

multiple regression analyses (Table 4). The addition of other site factors, namely precipitation, age or LAI improved the multiple regressions further. Generally, the goodness of fit between cDOY and NEP_c increased with the timing of later cDOYs, and with decreasing air temperatures (Table 4).

We conclude that lower mean annual temperatures lead to generally more pronounced winter net respiratory losses and it appears plausible that this is linked to later cDOYs. This is also in line with studies analyzing the onset of forests as a C-sink in relation to winter and spring temperatures [D D Baldocchi *et al.*, 2005; B I Cook *et al.*, 2012; Delpierre *et al.*, 2009; Monson *et al.*, 2011b]. Or the other way around, the warmer the site the less distinct the carbon loss period may be the earlier cDOY happens and the less likely the influence of cDOY on annual carbon uptake. Furthermore, we conclude that latter cDOYs are linked to lower annual NEP_c and, thus, the influence of cDOY on the annual NEP_c increases with its timing.

cDOYs of deciduous broad-leaf forests (DBF) showed the highest prediction quality for NEP_c (Table 4e) despite the fact that the respective cDOY did not correlate with mean annual temperature (Fig. 3) nor other site variables like for other forests types (Table 4d) or the pooled data (Table 4 a,b,c). Sites at higher altitudes (e.g. US-NR1, US-Me2) experience large inter-annual variation in the physiological active period, for example, 45 days at US-Me2 [Thomas *et al.* 2009] and studies in the mountains of the western US have shown declining snowpack for decades and its correlation with warm temperature anomalies. Further at US-NR1, longer growing seasons were correlated with low snow water equivalent and resulted in less annual net carbon uptake [Hu *et al.*, 2010]. Overall, such processes may confound an explanatory power of MAT for cDOY and NEP_c in certain cases, however we found no generally convincing explanation for the relationship between cDOY (its not found drivers) and NEP_c . Even when not understanding why DBF sites appear as a special case, we conclude that cDOY timing must in general depend on variables (eventually beyond the ones we analysed) containing information about the site and its past (climatic) history, including genetic predispositions leading to this high predictive power for NEP_c .

In summary, there are many indications for winter effects on NEP_c of forests and related to it on the cDOY timing. The compensation day (cDOY) is suggested to capture air temperature and intrinsic forest type-dependent differences, leading to a specific date in the first part of the calendar year, with a high explanatory power for the upcoming annual NEP_c values of the entire year for forest sites under distinct respiratory net carbon losses during wintertime.

4.5 Starting the NEP_c year in autumn

Three different ways of integrating NEP over a year were applied: the static ‘classical’ calendar-year method (January 1 to December 31), the static ‘Urbanski’ method (October 28 to October 27) and the more process-oriented ‘dynamic’ method, defining the ‘biological’ year as the period between two annual NEP_c peaks. There appeared distinctly better fits between cDOY and NEP_c for the two methods starting the NEP_c -year in autumn (Table 4). The classical method performed generally worse for all types of analyses (Tables 3, 4; Fig. 5). The additional gain of predictive quality for the dynamic method over the static Urbanski method was relatively small. This means that it is important to include the complete autumn and winter period before the actual C-sink period for interpreting NEP_c , but doing so with a static approach captures more or less the same information as when doing so with the site- and year-specific dynamic method (which can be more labor-intensive to deal with).

4.6 Conclusions

The compensation day cDOY reflects processes, which take place before the net C-sink period begins in forests in spring and early summer. The fact that cDOY explains more of NEP_c when starting the NEP_c -year in autumn shows that the (autumn-winter) period already before January 1st plays an important role for the following NEP_c performance. cDOY analysis takes seasonal and inter-annual variations of the carbon cycle dynamics into account, and is therefore suggested to take up carry-over effects of climate and carbon storage in temperate forests [Keenan *et al.*, 2012; Urbanski *et al.*, 2007; Zielis *et al.*, 2014; Zweifel *et al.*, 2010]. Such carry-over effects seem to be less important in forests with no distinct winter net respiratory loss of C (C-loss less than 10% of annual NEP_c). This is in line with the finding that cDOY gains explanatory power for NEP_c at sites with distinct winter respiratory C losses. The fact that biological processes, occurring before the annual net assimilation period begins, are able to explain more than 50% of the annual NEP_c , should change our view

on the drivers of NEP_c . Weather conditions during the main assimilation period of a forest seem to tell only half the story of the annual NEP_c . Thus, an accurate NEP_c interpretation additionally needs to include the conditions that affected a forest before this period.

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References

- Aubinet, M., T. Vesala, and D. Papale (2012), *Eddy covariance --- A Practical Guide to Measurement and Data Analysis*, 59-171 pp., Springer.
- Aurela, M., T. Laurila, and J. Tuovinen (2004), The timing of snow melt controls the annual CO₂ balance in a subarctic fen, *Geophysical Research Letters*, 31(16), L16119, doi:10.1029/2004GL020315.
- Baldocchi, D., and K. Wilson (2001), Modelling CO₂ and water vapour exchange of a temperate broadleaved forest across hourly to decadal time scales, *Ecological Modelling*, 142, 155-184.
- Baldocchi, D. D., T. Black, P. Curtis, E. Falge, J. Fuentes, A. Granier, L. Gu, A. Knohl, K. Pilegaard, and H. Schmid (2005), Predicting the onset of net carbon uptake by deciduous forests with soil temperature and climate data: a synthesis of FLUXNET data, *International journal of biometeorology*, 49(6), 377-387.
- Basler, D., and C. Körner (2014), Photoperiod and temperature responses of bud swelling and bud burst in four temperate forest tree species, *Tree physiology*, 34(4), 377-388, doi:10.1093/treephys/tpu021.
- Beringer, J., et al. (2016), An introduction to the Australian and New Zealand flux tower network - OzFlux, *Biogeosciences Discussions*, submitted.
- Brzostek, E. R., D. Dragoni, H. P. Schmid, A. F. Rahman, D. Sims, C. A. Wayson, D. J. Johnson, and R. P. Phillips (2014), Chronic water stress reduces tree growth and the carbon sink of deciduous hardwood forests, *Global change biology*, 20(8), 2531-2539.
- Campioli, M., A. Michelsen, A. Demey, A. Vermeulen, R. Samson, and R. Lemeur (2009), Net Primary Production and Carbon Stocks for Subarctic Mesic-Dry Tundras with Contrasting Microtopography, Altitude, and Dominant Species, *Ecosystems*, 12(5), 760-776.
- Churkina, G., D. Schimel, B. H. Braswell, and X. Xiao (2005), Spatial analysis of growing season length control over net ecosystem exchange, *Global Change Biology*, 11(10), 1777-1787, doi:10.1111/j.1365-2486.2005.001012.x.
- Ciais, P., et al. (2005), Europe-wide reduction in primary productivity caused by the heat and drought in 2003, *Nature*, 437(7058), 529-533, doi:10.1038/nature03972.
- Cook, B. D., K. J. Davis, W. Wang, A. Desai, B. W. Berger, R. M. Teclaw, J. G. Martin, P. V. Bolstad, P. S. Bakwin, and C. Yi (2004), Carbon exchange and venting anomalies in an upland deciduous forest in northern Wisconsin, USA, *Agricultural and Forest Meteorology*, 126(3), 271-295.
- Cook, B. I., E. M. Wolkovich, and C. Parmesan (2012), Divergent responses to spring and winter warming drive community level flowering trends, *Proceedings of the National Academy of Sciences*, 109(23), 9000-9005.

- Coursolle, C., H. Margolis, M.-A. Giasson, P.-Y. Bernier, B. Amiro, M. Arain, A. Barr, T. Black, M. Goulden, and J. McCaughey (2012), Influence of stand age on the magnitude and seasonality of carbon fluxes in Canadian forests, *Agricultural and Forest Meteorology*, *165*, 136-148.
- Delpierre, N., K. Soudani, C. Francois, B. Köstner, J. Y. PONTAILLER, E. Nikinmaa, L. Misson, M. Aubinet, C. Bernhofer, and A. Granier (2009), Exceptional carbon uptake in European forests during the warm spring of 2007: a data-model analysis, *Global Change Biology*, *15*(6), 1455-1474.
- Dragoni, D., H. P. Schmid, C. A. Wayson, H. Potter, C. S. B. Grimmond, and J. C. Randolph (2011), Evidence of increased net ecosystem productivity associated with a longer vegetated season in a deciduous forest in south-central Indiana, USA, *Global Change Biology*, *17*(2), 886-897.
- Epron, D., R. Liozon, and M. Mousseau (1996), Effects of elevated CO₂ concentration on leaf characteristics and photosynthetic capacity of beech (*Fagus sylvatica*) during the growing season, *Tree Physiology*, *16*(4), 425-432, doi:10.1093/treephys/16.4.425.
- European Fluxes Database Cluster (2014), Online resource: <http://www.europe-fluxdata.eu/>, last accessed on April 13, 2014.
- Flechard, C. R., et al. (2011), Dry deposition of reactive nitrogen to European ecosystems: a comparison of inferential models across the NitroEurope network, *Atmospheric Chemistry and Physics*, *11*(6), 2703-2728, doi:DOI 10.5194/acp-11-2703-2011.
- Gökkaya, K., V. Thomas, T. Noland, H. McCaughey, and P. Treitz (2014), Testing the robustness of predictive models for chlorophyll generated from spaceborne imaging spectroscopy data for a mixedwood boreal forest canopy, *International Journal of Remote Sensing*, *35*(1), 218-233.
- Gökkaya, K., V. Thomas, T. L. Noland, H. McCaughey, I. Morrison, and P. Treitz (2015), Prediction of macronutrients at the canopy level using spaceborne imaging spectroscopy and LiDAR data in a mixedwood boreal forest, *Remote Sensing*, *7*(7), 9045-9069.
- Gonsamo, A., J. M. Chen, D. T. Price, W. A. Kurz, and C. Wu (2012a), Land surface phenology from optical satellite measurement and CO₂ eddy covariance technique, *Journal of Geophysical Research: Biogeosciences*, *117*(G3).
- Gonsamo, A., J. M. Chen, C. Wu, and D. Dragoni (2012b), Predicting deciduous forest carbon uptake phenology by upscaling FLUXNET measurements using remote sensing data, *Agricultural and forest meteorology*, *165*, 127-135.
- Gough, C., C. Vogel, H. Schmid, H.-B. Su, and P. Curtis (2008), Multi-year convergence of biometric and meteorological estimates of forest carbon storage, *Agricultural and Forest Meteorology*, *148*(2), 158-170.

Gough, C. M., B. S. Hardiman, L. E. Nave, G. Bohrer, K. D. Maurer, C. S. Vogel, K. J. Nadelhoffer, and P. S. Curtis (2013), Sustained carbon uptake and storage following moderate disturbance in a Great Lakes forest, *Ecological Applications*, 23(5), 1202-1215.

Grünwald, T., and C. Bernhofer (2007), A decade of carbon, water and energy flux measurements of an old spruce forest at the Anchor Station Tharandt, *Tellus B*, 59(3), 387-396.

Herbst, M., M. Mund, R. Tamrakar, and A. Knohl (2015), Differences in carbon uptake and water use between a managed and an unmanaged beech forest in central Germany, *Forest Ecology and Management*, 355, 101-108.

Hoch, G., A. Richter, and C. Körner (2003), Non-structural carbon compounds in temperate forest trees, *Plant, Cell & Environment*, 26(7), 1067-1081.

Holmsgaard, E. (1955), Arringanalyser af danske skovtraeer, *Det forstl. forsøgsvaesen i Danmark*, 22(1).

Hu, J., D. J. Moore, S. P. Burns, and R. K. Monson (2010), Longer growing seasons lead to less carbon sequestration by a subalpine forest, *Global Change Biology*, 16(2), 771-783.

Jurik, T. W. (1986), Seasonal Patterns of Leaf Photosynthetic Capacity in Successional Northern Hardwood Tree Species, *American Journal of Botany*, 73(1), 131-138, doi:10.2307/2444285.

Keenan, T., et al. (2012), Terrestrial biosphere model performance for inter-annual variability of land-atmosphere CO₂ exchange, *Global Change Biology*, 18(6), 1971-1987, doi:10.1111/j.1365-2486.2012.02678.x.

Keith, H., E. Van Gorsel, K. L. Jacobsen, and H. A. Cleugh (2012), Dynamics of carbon exchange in a Eucalyptus forest in response to interacting disturbance factors, *Agricultural and Forest Meteorology*, 153, 67-81.

Knohl, A., E.-D. Schulze, O. Kolle, and N. Buchmann (2003), Large carbon uptake by an unmanaged 250-year-old deciduous forest in Central Germany, *Agricultural and Forest Meteorology*, 118(3), 151-167.

Koike, T. (1990), Autumn colouring, photosynthetic performance and leaf development of deciduous broadleaved trees in relation to forest succession, *Tree Physiology*, 7(1-4), 21-32.

Lagergren, F., A. Lindroth, E. Dellwik, A. Ibrom, H. Lankreijer, S. Launiainen, M. Mölder, P. Kolari, K. I. M. Pilegaard, and T. Vesala (2008), Biophysical controls on CO₂ fluxes of three Northern forests based on long-term eddy covariance data, *Tellus B*, 60(2), 143-152, doi:10.1111/j.1600-0889.2006.00324.x.

Law, B., E. Falge, L. v. Gu, D. Baldocchi, P. Bakwin, P. Berbigier, K. Davis, A. Dolman, M. Falk, and J. Fuentes (2002), Environmental controls over carbon dioxide and water vapor exchange of terrestrial vegetation, *Agricultural and Forest Meteorology*, 113(1), 97-120.

Law, B., and R. Waring (2015), Carbon implications of current and future effects of drought, fire and management on Pacific Northwest forests, *Forest Ecology and Management*, 355, 4-14.

Le Maire, G., N. Delapierre, M. Jung, P. Ciais, M. Reichstein, N. Viovy, A. Granier, A. Ibrom, P. Kolari, and B. Longdoz (2010), Detecting the critical periods that underpin interannual fluctuations in the carbon balance of European forests, *Journal of Geophysical Research: Biogeosciences* (2005–2012), 115(G3).

Lundmark, T., J. Heden, and J. E. Hallgren (1988), Recovery from winter depression of photosynthesis in pine and spruce, *Trees*, 2(2), 110-114, doi:10.1007/BF00196757.

McCaughey, J., M. Pejam, M. Arain, and D. Cameron (2006), Carbon dioxide and energy fluxes from a boreal mixedwood forest ecosystem in Ontario, Canada, *Agricultural and Forest Meteorology*, 140(1), 79-96.

Monson, R., D. Moore, N. Trahan, L. Scott-Denton, S. Burns, J. Hu, and D. Bowling (2011a), Process coupling and control over the response of net ecosystem CO₂ exchange to climate variability and insect disturbance in subalpine forests of the Western US, paper presented at AGU Fall Meeting Abstracts, No. B13J-01.

Monson, R., D. Moore, N. Trahan, L. Scott-Denton, S. Burns, J. Hu, and D. Bowling (2011b), Process coupling and control over the response of net ecosystem eCO₂ exchange to climate variability and insect disturbance in subalpine forests of the Western US, paper presented at AGU Fall Meeting Abstracts, No. B13J-01.

Mund, M., W. Kutsch, C. Wirth, T. Kahl, A. Knohl, M. Skomarkova, and E.-D. Schulze (2010), The influence of climate and fructification on the inter-annual variability of stem growth and net primary productivity in an old-growth, mixed beech forest, *Tree physiology*, 30(6), 689-704.

Munger, J. W., S.-M. Fan, P. S. Bakwin, M. L. Goulden, A. Goldstein, A. S. Colman, and S. C. Wofsy (1998), Regional budgets for nitrogen oxides from continental sources: Variations of rates for oxidation and deposition with season and distance from source regions, *Journal of Geophysical Research*, 103(D7).

Munger, J. W., S. C. Wofsy, P. S. Bakwin, S. M. Fan, M. L. Goulden, B. C. Daube, A. H. Goldstein, K. E. Moore, and D. R. Fitzjarrald (1996), Atmospheric deposition of reactive nitrogen oxides and ozone in a temperate deciduous forest and a subarctic woodland: 1. Measurements and mechanisms, *Journal of Geophysical Research: Atmospheres*, 101(D7), 12639-12657.

Nave, L. E., C. S. Vogel, C. M. Gough, and P. S. Curtis (2009), Contribution of atmospheric nitrogen deposition to net primary productivity in a northern hardwood forest, *Canadian journal of forest research*, 39(6), 1108-1118.

Ottander, C., D. Campbell, and G. Oquist (1995), Seasonal changes in photosystem II organization and pigment composition in *Pinus sylvestris*, *Planta*, 197(1), 176-183, doi:10.1007/BF00239954.

Ottander, C., and G. Oquist (1991), Recovery of Photosynthesis in Winter-Stressed Scots Pine, *Plant Cell Environ*, *14*(3), 345-349, doi:DOI 10.1111/j.1365-3040.1991.tb01511.x.

Pallardy, S. G. (2010), *Physiology of woody plants*, Academic Press.

Pilegaard, K., A. Ibrom, M. S. Courtney, P. Hummelshøj, and N. O. Jensen (2011), Increasing net CO₂ uptake by a Danish beech forest during the period from 1996 to 2009, *Agricultural and Forest Meteorology*, *151*(7), 934-946.

Pita, G., B. Gielen, D. Zona, A. Rodrigues, S. Rambal, I. A. Janssens, and R. Ceulemans (2013), Carbon and water vapor fluxes over four forests in two contrasting climatic zones, *Agricultural and forest meteorology*, *180*, 211-224.

Quinn, G., and M. Keough (2002), *Experimental design and data analysis for biologists*, Cambridge University Press.

R Development Core Team (2013), R: A language and environment for statistical computing: <http://www.R-project.org>, Vienna, Austria: R Foundation for Statistical Computing, 1-1731.

Reich, P. B., M. B. Walters, and D. S. Ellsworth (1991), Leaf Age and Season Influence the Relationships between Leaf Nitrogen, Leaf Mass Per Area and Photosynthesis in Maple and Oak Trees, *Plant Cell Environ*, *14*(3), 251-259, doi:DOI 10.1111/j.1365-3040.1991.tb01499.x.

Reichstein, M., et al. (2013), Climate extremes and the carbon cycle, *Nature*, *500*(7462), 287-295, doi:10.1038/nature12350.

Reichstein, M., et al. (2005), On the separation of net ecosystem exchange into assimilation and ecosystem respiration: review and improved algorithm, *Global Change Biology*, *11*(9), 1424-1439, doi:DOI 10.1111/j.1365-2486.2005.001002.x.

Richardson, A. D., et al. (2010), Influence of spring and autumn phenological transitions on forest ecosystem productivity, *Philosophical transactions of the Royal Society of London. Series B, Biological sciences*, *365*(1555), 3227-3246, doi:10.1098/rstb.2010.0102.

Richardson, A. D., D. Y. Hollinger, D. B. Dail, J. T. Lee, J. W. Munger, and J. O'Keefe (2009), Influence of spring phenology on seasonal and annual carbon balance in two contrasting New England forests, *Tree Physiol*, *29*(3), 321-331, doi:10.1093/treephys/tpn040.

Rodrigues, A., G. Pita, J. Mateus, C. Kurz-Besson, M. Casquilho, S. Cerasoli, A. Gomes, and J. Pereira (2011), Eight years of continuous carbon fluxes measurements in a Portuguese eucalypt stand under two main events: drought and felling, *Agricultural and Forest Meteorology*, *151*(4), 493-507.

Roman, D., K. Novick, E. Brzostek, D. Dragoni, F. Rahman, and R. Phillips (2015), The role of isohydric and anisohydric species in determining ecosystem-scale response to severe drought, *Oecologia*, *179*(3), 641-654.

Scartazza, A., D. Di Baccio, P. Bertolotto, O. Gavrichkova, and G. Matteucci (2016), Investigating the European beech (*Fagus sylvatica* L.) leaf characteristics along the vertical

canopy profile: leaf structure, photosynthetic capacity, light energy dissipation and photoprotection mechanisms, *Tree physiology*, tpw038.

Schmid, H. P., C. S. B. Grimmond, F. Cropley, B. Offerle, and H.-B. Su (2000), Measurements of CO₂ and energy fluxes over a mixed hardwood forest in the mid-western United States, *Agricultural and Forest Meteorology*, 103(4), 357-374.

Schwarz, P., B. Law, M. Williams, J. Irvine, M. Kurpius, and D. Moore (2004), Climatic versus biotic constraints on carbon and water fluxes in seasonally drought-affected ponderosa pine ecosystems, *Global Biogeochemical Cycles*, 18(4).

Shao, J., X. Zhou, Y. Luo, B. Li, M. Aurela, D. Billesbach, P. D. Blanken, R. Bracho, J. Chen, and M. Fischer (2016), Direct and indirect effects of climatic variations on the interannual variability in net ecosystem exchange across terrestrial ecosystems, *Tellus B*, 68.

Sievering, H., T. Kelly, G. McConville, C. Seibold, and A. Turnipseed (2001), Nitric acid dry deposition to conifer forests:: Niwot Ridge spruce–fir–pine study, *Atmospheric environment*, 35(22), 3851-3859.

Suni, T., F. Berninger, T. Vesala, T. Markkanen, P. Hari, A. Mäkelä, H. Ilvesniemi, H. Hänninen, E. Nikinmaa, and T. Huttula (2003), Air temperature triggers the recovery of evergreen boreal forest photosynthesis in spring, *Global Change Biology*, 9(10), 1410-1426.

Thomas, C. K., B. E. Law, J. Irvine, J. G. Martin, J. C. Pettijohn, and K. J. Davis (2009), Seasonal hydrology explains interannual and seasonal variation in carbon and water exchange in a semiarid mature ponderosa pine forest in central Oregon, *Journal of Geophysical Research: Biogeosciences*, 114(G4).

Urbanski, S., C. Barford, S. Wofsy, C. Kucharik, E. Pyle, J. Budney, K. McKain, D. Fitzjarrald, M. Czikowsky, and J. Munger (2007), Factors controlling CO₂ exchange on time scales from hourly to decadal at Harvard Forest, *Journal of Geophysical Research: Biogeosciences*, 112(G2), 1-25 G02020, doi:10.1029/2006JG000293.

Wolf, S., W. Eugster, C. Ammann, M. Hani, S. Zielis, R. Hiller, J. Stieger, D. Imer, L. Merbold, and N. Buchmann (2013), Contrasting response of grassland versus forest carbon and water fluxes to spring drought in Switzerland, *Environmental Research Letters*, 8(3), 035007, doi:doi 10.1088/1748-9326/8/3/035007.

Wu, C., et al. (2013), Inter-annual variability of net ecosystem productivity in forests is explained by carbon flux phenology in autumn, *Global Ecology and Biogeography*, 22, 994-1006.

Wu, J., L. van der Linden, G. Lasslop, N. Carvalhais, K. Pilegaard, C. Beier, and A. Ibrom (2012), Effects of climate variability and functional changes on the inter-annual variation of the carbon balance in a temperate deciduous forest, *Biogeosciences*, 9(1), 13-28.

Zielis, S., S. Etzold, R. Zweifel, W. Eugster, M. Haeni, and N. Buchmann (2014), NEP of a Swiss subalpine forest is significantly driven not only by current but also by previous year's weather, *Biogeosciences*, 11, 1627-1635.

Zweifel, R., W. Eugster, S. Etzold, M. Dobbertin, N. Buchmann, and R. Hasler (2010), Link between continuous stem radius changes and net ecosystem productivity of a subalpine Norway spruce forest in the Swiss Alps, *New Phytologist*, 187(3), 819-830, doi:DOI 10.1111/j.1469-8137.2010.03301.x.

Zweifel, R., and R. Hasler (2000), Frost-induced reversible shrinkage of bark of mature subalpine conifers, *Agricultural and Forest Meteorology*, 102(4), 213-222, doi:Doi 10.1016/S0168-1923(00)00135-0.

Zweifel, R., H. Item, and R. Hasler (2000), Stem radius changes and their relation to stored water in stems of young Norway spruce trees, *Trees-Struct Funct*, 15(1), 50-57, doi:DOI 10.1007/s004680000072.

Zweifel, R., L. Zimmermann, F. Zeugin, and D. M. Newbery (2006), Intra-annual radial growth and water relations of trees: implications towards a growth mechanism, *J Exp Bot*, 57(6), 1445-1459, doi:10.1093/jxb/erj125.

Table 1. Selected characteristics of the study sites (see also Fig. 2). Site names are abbreviated with the FLUXNET codes. Forest types are abbreviated as: DBF = deciduous broad-leaf forests, ENF = evergreen needle-leaf forests, MF = mixed forests, and EBF = evergreen broad-leaf forests. Further terms in the header have the following meaning: Lat. = Latitude (degrees, WGS84); Long = longitude (degrees, WGS84) of the site. Years used = data of the indicated years used in this study. MAP = mean annual sum of precipitation (mm). MAT = mean annual air temperature (°C) at the top of the eddy towers; Altitude = meter above sea level (m asl); Age = average age of the mature trees in the stand (years); Height = maximum tree height (m); LAI = leaf area index (m m⁻²); N = mean annual nitrogen deposition (kg ha⁻¹ yr⁻¹), most data and where stated from *Flechard et al.* [2011]; Data access: NEP_c data were obtained from European Database Cluster in Level 4 quality, from the FluxNet2015 dataset, or from the respective PIs incl. own sites (see Material and Method). Site characteristics were obtained from <https://fluxnet.ornl.gov>, from the *FluxNet2015 Metadata excel sheet, from <http://www.bgc-jena.mpg.de/public/carboeur/sites/SITE.html> or from *personal communication (pc). Not available data are indicated with “–”.

Code	Name	Forest Type	Lat.	Long.	Years	MAP	MAT	Altitude	Age	Height	LAI	N	Data access	Data in this table from...
					Used									
AU-TUM	Tumbarumba	EBF	-35.66	148.15	2002-2013	1924	9.6	1249	100	40	2.5	--	FluxNet2015	Beringer et al., 2016; FluxNet2015 meta*; pc*
BE-VIE	Vielsalm	MF	50.31		6 1997-2014	1062	7.8	493	95	35	4.6	10.2	FluxNet2015	Flechard et al., 2011; FluxNet2015 meta; pc
CA-GRO	Ontario, Groundhog River, Boreal Mixedwood Forest	MF	48.22	-82.16	2004-2013	831	1.3	340	84	32	--	--	FluxNet2015	McCaughey et al., 2006; Gökkaya et al., 2014; Gökkaya et al., 2015; FluxNet2015 meta; pc
CA-QFO	Quebec, Eastern Boreal, Mature Black Spruce	ENF	49.69	-74.34	2004-2010	962	-0.36	382	105	13.8	3.7	--	FluxNet2015	Coursolle et al., 2012; FluxNet2015 meta; pc
CH-DAV	Davos	ENF	46.81	9.86	1997-2011	1046	3.5	1662	240	25	3.9	1.5	own site	own site; FluxNet2015 meta
CH-LAE	Lägeren	MF	47.48	8.37	2004-2011	1211	8.7	682	140	31	3.6	14.3	own site	own site; Flechard et al., 2011; FluxNet2015 meta
CZ-BK1	Bily Kriz	ENF	49.5	18.54	2001-2010	1316	6.7	875	30	10	7.5	10.5	download, L4	Flechard et al., 2011; pc
DE-HAI	Hainich	DBF	51.08	10.45	2000-2012	720	8.3	430	125	33	6	12.6	FluxNet2015	Knohl et al., 2003; Mund et al., 2010; Flechard et al., 2011; Herbst et al., 2015; FluxNet2015 meta; pc
DE-THA	Tharandt	ENF	50.96	13.57	1997-2014	820	7.7	385	125	26.5	7.6	12.5	FluxNet2015	Grünwald and Bernhofer, 2007; Flechard et al., 2011; FluxNet2015 meta; pc
DK-SOR	Soroe	DBF	55.49	11.65	1997-2014	660	8.2	40	100	35	4.8	14.6	FluxNet2015	Pilegaard et al., 2011; FluxNet2015 meta; pc
ES-ES1	El Saler	ENF	39.35	-0.32	2000-2007	551	17.9	1	100	10	3.5	16.9	download, L4	FluxNet2015 meta; pc
FI-HYY	Hyytiälä	ENF	61.85	24.3	1997-2014	709	3.8	181	40	14	2.5	3.5	download, L4	Flechard et al., 2011; FluxNet2015 meta; pc
FR-PUE	Puechabon	EBF	43.74	3.6	2000-2014	883	13.5	270	73	5.5	2.9	--	FluxNet2015	Flechard et al., 2011; FluxNet2015 meta; pc
IT-COL	Collelongo	DBF	41.85	13.59	1997-2011	1180	6.3	1645	105	25	5	5.3	download, L4	Scartazza et al., 2016; Flechard et al., 2011; FluxNet2015 meta; pc
IT-LAV	Lavarone	ENF	45.96	11.28	2003-2014	1291	7.8	1353	100	28	8.1	15.4	FluxNet2015	FluxNet2015 meta; pc
IT-REN	Renon	ENF	46.59	11.43	1999-2013	809	4.7	1730	200	31	5.1	4.8	FluxNet2015	Flechard et al., 2011; FluxNet2015 meta; pc
NL-LOO	Loobos	ENF	52.17	5.74	1997-2013	966	10	25	90	18	1.9	32.4	FluxNet2015	Flechard et al., 2011; FluxNet2015 meta; pc
PL-TUC	Tuczno	ENF	53.21	16.1	2008-2011	625	7.8	105	54	20	1.1	8.5	STSM, direct	FluxNet2015 meta; pc
PT-ESP	Espirra	EBF	38.64	-8.6	2002-2008	665	15.4	85	12	20	3.1	7.2	download, L4	Pita et al., 2013; Rodrigues et al., 2011; Flechard et al., 2011; pc
RU-FYO	Fyodorovskoye	ENF	56.46	32.92	1998-2014	711	3.9	265	196	25.7	3.5	8.3	FluxNet2015	Flechard et al., 2011; FluxNet2015 meta; pc
SE-NOR	Norunda	ENF	60.09	16.22	1996-2011	527	5.5	45	100	25	5	3.4	STSM, direct	Lagergren et al., 2008; Flechard et al., 2011; pc
US-HA1	Harvard Forest EMS Tower (HFR1)	DBF	42.54	-72.15	1992-2012	1071	6.62	340	80	18	3.7	6.4	FluxNet2015	Munger et al., 1996; Munger et al., 1998; FluxNet2015 meta, pc
US-ME2	Metolius-intermediate aged ponderosa pine	ENF	44.45	-121.55	2002-2014	523	6.28	1253	90	14	3.2	1	FluxNet2015	Schwarz et al., 2004; Thomas et al., 2009; FluxNet2015 meta; pc
US-MMS	Morgan Monroe State Forest	DBF	39.92	-86.41	1999-2014	1032	10.9	275	90	27	4.7	6	FluxNet2015	Schmid et al., 2000; Dragoni et al., 2011; Brzostek et al., 2014; Roman et al., 2015; FluxNet2015 meta; pc
US-NR1	Niwot Ridge Forest (LTER NWT1)	ENF	40.03	-105.55	1999-2014	800	1.5	3050	115	13	4	6	FluxNet2015	Sievering et al., 2001; FluxNet2015 meta; pc
US-UMB	Univ. of Mich. Biological Station	DBF	45.56	-84.71	2000-2014	803	5.38	234	93	22	3.5	7.5	FluxNet2015	Gough et al., 2008; Nave et al., 2009; C M Gough et al., 2013; FluxNet2015 meta; pc
US-WCR	Willow Creek	DBF	45.8	-90.08	1999-2014	787	4.02	520	70	24.2	5.36	--	FluxNet2015	B D Cook et al., 2004; FluxNet2015 meta; pc

References in Tab 1.: [Beringer et al., 2016; Brzostek et al., 2014; B D Cook et al., 2004; Coursolle et al., 2012; Dragoni et al., 2011; Flechard et al., 2011; Gökkaya et al., 2014; Gökkaya et al., 2015; C Gough et al., 2008; C M Gough et al., 2013; Grünwald and Bernhofer, 2007; Herbst et al., 2015; Keith et al., 2012; Knohl et al., 2003; Lagergren et al., 2008; McCaughey et al., 2006; Mund et al., 2010; Munger et al., 1998; Munger et al., 1996; Nave et al., 2009; Pilegaard et al., 2011; Pita et al., 2013; Rodrigues et al., 2011; Roman et al., 2015; Scartazza et al., 2016; Schmid et al., 2000; Schwarz et al., 2004; Sievering et al., 2001; Thomas et al., 2009]

Table 2. NEP_c averages, cDOY averages, and relationships between compensation days (cDOY = day of year when winter respiratory losses were compensated) and integrated net ecosystem productivity (NEP_c) for each site and integration method (classical, Urbanski, dynamic). Mean NEP_c = mean annual net ecosystem productivity and standard deviation (g m⁻² yr⁻¹), averaged over all three integration methods (classical, Urbanski, dynamic); n cDOYs = how many cDOYs could be calculated for each of the integration methods; Mean cDOY = mean NEP_c compensation day of the years investigated incl. standard deviations. Average SD NEP = standard deviation of NEP_c averaged over all three integration methods; % winter respiratory loss = average percentage of winter respiratory loss from average NEP_c; Classification into whether the site is a site with or without distinct winter respiratory loss (threshold is <10%).

Site	Mean NEP _c	n cDOY cla.	n cDOY Urb.	n cDOY dyn.	cDOY	R2 cla.	R2 cla.	p cla.	R2 Urb.	R2 Urb.	p Urb.	R2 dyn.	R2 dyn.	p dyn.	Average SD NEP _c	% winter resp. Loss	Classification
AU-TUM	1006.7 ±273.4	12	--	12	Jul 03 ±6	0.05		0.48	--		--	0.05		0.51	159.9	5.4%	without
BE-VIE	461.9 ±155.6	18	18	18	Apr 29 ±28	0.62		0.00	0.73		0.00	0.74		0.00	173.3	-19.8%	with
CA-GRO	120.6 ±47.3	10	10	10	Jun 27 ±15	0.40		0.05	0.52		0.02	0.70		0.00	52.8	-143.6%	with
CA-QFO	2.8 ±14.5	7	7	3	Jul 23 ±23	0.08		0.54	0.24		0.26	0.01		0.93	14.7	-420.3%	with
CH-DAV	134.1 ±70.8	15	15	15	Mai 28 ±34	0.74		0.00	0.88		0.00	0.87		0.00	74.0	-80.8%	with
CH-LAE	579.7 ±124.5	8	7	8	Mai 10 ±14	0.76		0.00	0.72		0.02	0.68		0.01	122.2	-16.1%	with
CZ-BK1	798.9 ±132.8	10	10	10	Apr 04 ±8	0.07		0.45	0.00		0.94	0.00		0.95	128.9	-6.6%	without
DE-HAI	579.2 ±73.2	13	13	13	Jun 21 ±6	0.12		0.25	0.12		0.25	0.06		0.42	68.0	-51.4%	with
DE-THA	617.9 ±81.1	18	18	18	Apr 01 ±18	0.16		0.10	0.18		0.08	0.20		0.07	81.7	-7.5%	without
DK-SOR	176 ±139.4	18	18	16	Jun 26 ±18	0.66		0.00	0.84		0.00	0.82		0.00	151.7	-88.6%	with
ES-ES1	415.5 ±154.9	8	6	8	Jan 27 ±27	0.06		0.57	0.01		0.85	0.47		0.06	164.9	-3.6%	without
FI-HYY	244.9 ±50.5	16	16	16	Mai 26 ±10	0.19		0.10	0.53		0.00	0.59		0.00	53.3	-41.6%	with
FR-PUE	230.5 ±89.2	14	12	14	Feb 01 ±22	0.01		0.72	0.19		0.16	0.06		0.41	88.1	-8.6%	without
IT-COL	622 ±179	13	13	13	Jun 09 ±10	0.45		0.01	0.53		0.00	0.53		0.00	179.8	-28.2%	with
IT-REN	248 ±96.7	12	12	12	Mai 10 ±28	0.11		0.29	0.04		0.52	0.03		0.61	85.9	-7.4%	without
NL-LOO	427.5 ±144.6	17	16	17	Mrz 22 ±41	0.52		0.00	0.60		0.00	0.70		0.00	160.7	-9.7%	without
PL-TUC	667.5 ±76.1	4	1	4	Jan 28 ±30	0.53		0.28	0.00		--	0.70		0.16	82.9	0.1%	without
PT-ESP	459.7 ±356.6	6	2	6	Feb 02 ±50	0.48		0.13	1.00		--	0.51		0.11	363.0	-6.5%	without
RU-FYO	-136.7 ±128.2	11	8	2	Jun 13 ±26	0.33		0.07	0.19		0.29	1.00		--	145.4	104.7%	with
SE-NOR	52.8 ±33.8	15	15	12	Apr 15 ±31	0.67		0.00	0.86		0.00	0.51		0.01	37.1	-52.0%	with
US-HA1	211.2 ±188	21	19	19	Jul 24 ±22	0.54		0.00	0.75		0.00	0.78		0.00	102.5	-401.0%	with
US-ME2	588 ±291.2	13	5	13	Feb 08 ±40	0.59		0.00	0.76		0.05	0.67		0.00	316.6	-0.9%	without
US-MMS	428.6 ±75.8	16	16	16	Jun 20 ±9	0.15		0.14	0.19		0.09	0.24		0.05	39.3	-60.9%	with
US-NR1	170.7 ±30.1	16	16	16	Jun 02 ±9	0.01		0.71	0.09		0.25	0.12		0.19	31.8	-48.3%	with
US-UMB	268.9 ±65.8	15	15	15	Jul 04 ±10	0.19		0.10	0.27		0.05	0.34		0.02	31.4	-80.5%	with
US-WCR	271.9 ±153.6	12	12	11	Jun 29 ±41	0.35		0.04	0.55		0.01	0.67		0.00	164.6	-89.7%	with

Table 3. Stepwise multiple linear regression models to determine the drivers of the day of compensation cDOY for the a) classical, b) Urbanski and c) dynamic integration method. The variables were included one-by-one in the models (MAT = mean annual air temperature at the top of the eddy towers; LAI = leaf area index; Age = average age of the mature trees in the stand; Height = maximum tree height; N = mean annual nitrogen deposition; MAP = mean annual sum of precipitation). The β -coefficients (var) indicate the relative importance of the variable, ranging from -1 (highest importance, negative correlation) to +1 (highest importance, positive correlation). The first column gives the R^2 for individual site characteristics (see Table 1). *: $p < 0.05$, **: $p < 0.01$, ***: $p < 0.001$.

a) classical integration method

Drivers of cDOY	R2 alone	1 var	2 vars	3 vars	4 vars	5 vars	6 vars	7 vars
MAT	0.4***	-0.63	-0.58	-0.57	-0.58	-0.53	-0.55	-0.48
LAI	0.11	--	0.25	0.25	0.23	0.13	0.14	0.03
Age	0.06	--	--	0.03	-0.02	-0.06	-0.03	0.06
Height	0.05	--	--	--	0.1	0.29	0.27	0.23
N	0.03	--	--	--	--	0.08	0.07	-0.05
Altitude	0.01	--	--	--	--	--	-0.07	-0.19
MAP	0	--	--	--	--	--	--	0.37
Total adjR2	--	0.4***	0.39***	0.36**	0.34**	0.17*	0.12*	0.21*

b) Urbanski integration method

Drivers of cDOY	R2 alone	1 var	2 vars	3 vars	4 vars	5 vars	6 vars	7 vars
MAT	0.45***	-0.67	-0.61	-0.58	-0.59	-0.61	-0.43	-0.5
Height	0.17*	--	0.29	0.29	0.31	0.34	0.35	0.27
MAP	0.1	--	--	0.21	0.21	0.21	0.28	0.33
Age	0.07	--	--	--	-0.06	-0.09	-0.04	0.06
LAI	0.06	--	--	--	--	-0.01	-0.03	-0.01
N	0.05	--	--	--	--	--	-0.06	-0.12
Altitude	0.01	--	--	--	--	--	--	-0.24
Total adjR2	--	0.45***	0.48***	0.51***	0.49***	0.45***	0.23*	0.21*

c) dynamic integration method

Drivers of cDOY	R2 alone	1 var	2 vars	3 vars	4 vars	5 vars	6 vars	7 vars
MAT	0.4***	-0.63	-0.57	-0.57	-0.57	-0.5	-0.43	-0.52
LAI	0.1	--	0.23	0.23	0.22	0.1	0.01	0.04
Age	0.06	--	--	0.03	0	-0.06	-0.03	0.1
Height	0.03	--	--	--	0.06	0.36	0.36	0.24
N	0.03	--	--	--	--	0.04	-0.02	-0.09
MAP	0.01	--	--	--	--	--	0.23	0.3
Altitude	0	--	--	--	--	--	--	-0.33
Total adjR2	--	0.4***	0.37**	0.34**	0.31**	0.2*	0.2*	0.22*

Table 4. Stepwise multiple linear regression models to determine the drivers of net ecosystem productivity NEP_c for the a) classical (all sites), b) Urbanski (all sites) and c) dynamic integration method (all sites). d) shows the analysis for the dynamic integration for evergreen needle-leaf forests (ENF) only, and e) shows the analysis for the dynamic integration for deciduous broad-leaf forests (DBF) only. Other forest types had too low replications ($n=3$) for a separate analysis. The variables were included one-by-one in the models (cDOY = compensation day; MAT = mean annual air temperature at the top of the eddy towers; LAI = leaf area index; Age = average age of the mature trees in the stand; Height = maximum tree height; N = mean annual nitrogen deposition; MAP = mean annual sum of precipitation). The β -coefficients (var) indicate the relative importance of the variable, ranging from -1 (highest importance, negative correlation) to +1 (highest importance, positive correlation). The first column gives the R^2 for individual site characteristics (see Table 1). *: $p<0.05$, **: $p<0.01$, ***: $p<0.001$.

a) classical integration method (all sites)

Drivers of NEP_c	R2 alone	1 var	2 vars	3 vars	4 vars	5 vars	6 vars	7 vars	8 vars
cDOY	0.23*	-0.48	-0.48	-0.39	-0.44	-0.38	-0.36	-0.59	-0.6
MAP	0.22*	--	0.47	0.47	0.41	0.33	0.26	0.45	0.38
MAT	0.14	--	--	0.14	0.14	0.27	0.25	0.14	0.1
Height	0.05	--	--	--	0.21	0.24	0.36	0.39	0.33
Altitude	0.04	--	--	--	--	0.24	0.38	0.38	0.35
Age	0.04	--	--	--	--	--	-0.31	-0.39	-0.38
N	0.02	--	--	--	--	--	--	0.04	0.07
LAI	0.02	--	--	--	--	--	--	--	0.21
Total adjR2	--	0.23*	0.45**	0.47**	0.51**	0.55**	0.61**	0.54	0.57

b) Urbanski integration method (all sites)

Drivers of NEP_c	R2 alone	1 var	2 vars	3 vars	4 vars	5 vars	6 vars	7 vars	8 vars
cDOY	0.35**	-0.59	-0.7	-0.74	-0.67	-0.64	-0.83	-0.85	-0.85
MAP	0.23*	--	0.38	0.38	0.33	0.31	0.34	0.42	0.37
MAT	0.15	--	--	-0.06	0.06	0.06	0	0.01	-0.01
Altitude	0.05	--	--	--	0.18	0.24	0.34	0.31	0.29
Age	0.04	--	--	--	--	-0.14	-0.33	-0.33	-0.33
Height	0.03	--	--	--	--	--	0.45	0.38	0.34
N	0.02	--	--	--	--	--	--	-0.03	-0.01
LAI	0	--	--	--	--	--	--	--	0.14
Total adjR2	--	0.35**	0.47**	0.48**	0.5**	0.51**	0.65***	0.69**	0.71**

c) dynamic integration method (all sites)

Drivers of NEP_c	R2 alone	1 var	2 vars	3 vars	4 vars	5 vars	6 vars	7 vars	8 vars
cDOY	0.47***	-0.68	0.39	-0.66	-0.6	-0.58	-0.6	-0.84	-0.86
MAP	0.22*	--	-0.64	0.39	0.35	0.34	0.23	0.45	0.39
MAT	0.14	--	--	-0.04	0.06	0.05	0.06	0	-0.04
Altitude	0.06	--	--	--	0.15	0.2	0.29	0.27	0.24
Age	0.04	--	--	--	--	-0.13	-0.25	-0.31	-0.3

Height	0.02	--	--	--	--	--	0.29	0.38	0.33
N	0.02	--	--	--	--	--	--	0.01	0.03
LAI	0.01	--	--	--	--	--	--	--	0.18
Total adjR2	--	0.47***	0.62**	0.62***	0.63***	0.65***	0.7***	0.72**	0.74**

d) dynamic integration method (ENF only)

Drivers of NEPc	R2 alone	1 var	2 vars	3 vars	4 vars	5 vars	6 vars	7 vars	8 vars
cDOY	0.58**	-0.76	-0.4	-1.03	-1.08	-1.16	-1.09	-1.04	-0.71
MAT	0.17	--	-1.07	-0.4	-0.42	-0.38	-0.36	-0.22	0.43
Age	0.13	--	--	-0.09	-0.11	-0.06	-0.04	-0.2	-0.72
LAI	0.06	--	--	--	0.35	0.2	0.23	0.15	-0.04
MAP	0.04	--	--	--	--	0.36	0.36	0.46	0.56
N	0.03	--	--	--	--	--	0	-0.1	-0.25
Height	0.01	--	--	--	--	--	--	0.22	0.71
Altitude	0.01	--	--	--	--	--	--	--	0.53
Total adjR2	--	0.58**	0.65	0.66**	0.78**	0.87***	0.87**	0.88*	0.94*

e) dynamic integration method (DBF only)

Drivers of NEPc	R2 alone	1 var	2 vars	3 vars	4 vars	5 vars	6 vars	7 vars	8 vars
cDOY	0.68*	-0.83	0.37	-0.44	-0.4	-0.37	--	--	--
Altitude	0.53	--	-0.61	0.43	0.44	0.16	--	--	--
Age	0.33	--	--	0.23	0.22	0.29	--	--	--
LAI	0.27	--	--	--	0.07	0.24	--	--	--
MAP	0.17	--	--	--	--	0.38	--	--	--
N	0.11	--	--	--	--	--	--	--	--
MAT	0.03	--	--	--	--	--	--	--	--
Height	0.02	--	--	--	--	--	--	--	--
Total adjR2	--	0.68*	0.78*	0.81*	0.81	0.86	--	--	--

Figure 1. Three different methods of integrating net ecosystem productivity (NEP_c) over time (real data shown: Hyytiälä, years 2010 to 2012): ‘Classical’ integration runs from Jan 1 to Dec 31 of each calendar year, ‘current’ is year 2; ‘Urbanski’ integration from Oct 28 of the previous year to the end of Oct 27 of the current year; ‘Dynamic’ integration runs dynamically for every site and for every year from the day of the previous year’s cumulated NEP_c maximum (MAX_{NEP_c}) to the current year’s cumulated NEP_c maximum ($MAX_{NEP_c}(+1)$). Net carbon losses occur between MAX_{NEP_c} and the minimum of NEP_c of the current year (MIN_{NEP_c}). The day of compensation (cDOY) is defined as the day of the year when MAX_{NEP_c} (of the previous year) is crossed by NEP_c in the current year. Accordingly, cDOY depends on the integration method. For the Southern hemisphere, i.e. for the Australian site AU-Tum, the same cuts were made, one half year later. The corresponding year started on July 1 and ended on June 30.

Figure 2. Spatial distribution of 25 sites across North America and Europe and one Australian site (not shown). Site abbreviations are listed in Table 1. ENF = evergreen needle-leaf forests, DBF = deciduous broad-leaf forests, MF = mixed forests, and EBF = evergreen broad-leaf forests.

Figure 3. Relationship between cDOY (integrated dynamically) and mean annual temperature (MAT) grouped for the four forest types. **a)** all sites included (n = 26) and **b)** Sites with distinct winter respiratory losses only, i.e., where winter net respiration loss accounts for more than 10% of the annual net ecosystem productivity and n was > 2 for all integration methods.

Figure 4. Linear regressions between compensation days (cDOY) and annual sums of net ecosystem productivity (NEP_c) (Table 1) for each site and integration method (a) classical, b) Urbanski and c) dynamic). Solid regression lines are shown for $R^2 \geq 0.4$, broken lines for the rest. The frequency columns at the bottom of each panel indicate the number of site years occurring at a specific cDOY (color-coded for the four forest types).

Figure 5. Differences between modelled annual net ecosystem productivity (NEP_c), as a linear function of cDOY and measured NEP_c for each site in a leave-one-year-out cross-evaluation for each integration method: **a)** classical, **b)** Urbanski and **c)** dynamic. The results are grouped for sites with and without distinct winter net respiratory losses. The Urbanski method was left out for AU-Tum, there, the ‘year’ begins on July 1 and goes to June 30 one year later.